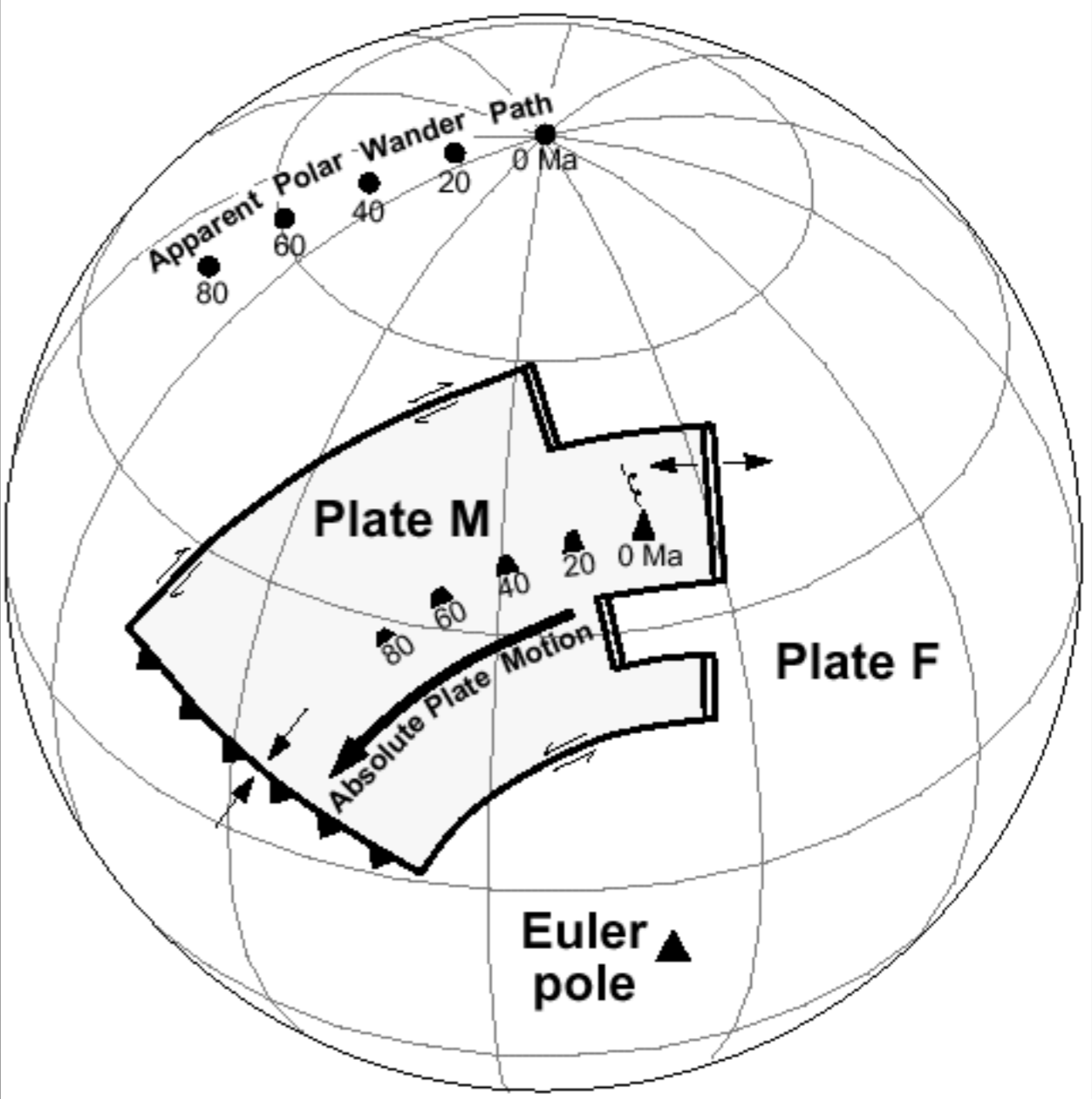




The Geologic Story of the Carpathians

Mihai Ducea
University of Arizona
Department of Geosciences,
Tucson, AZ 85718









Puerto Vallarta



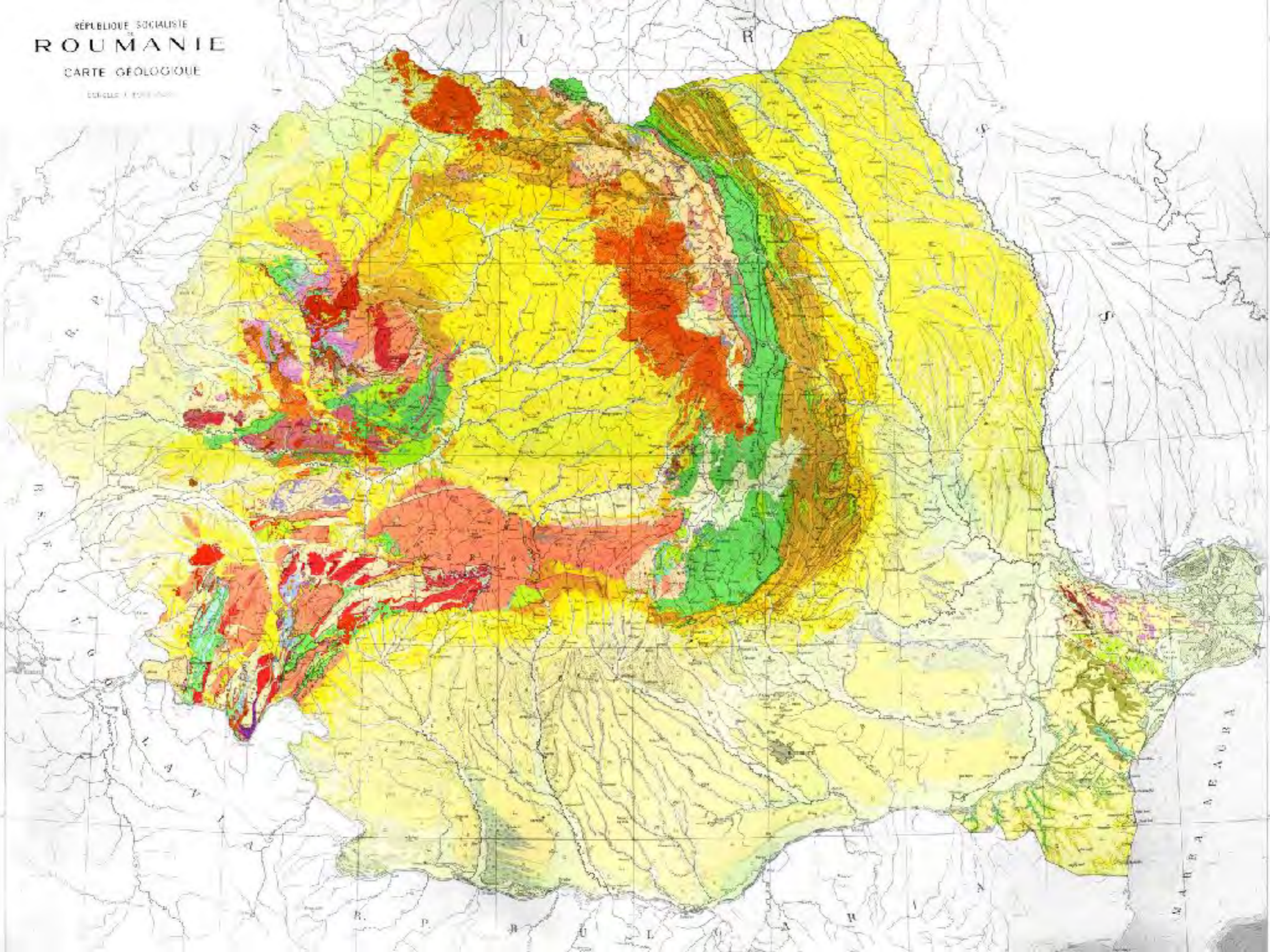
Summary

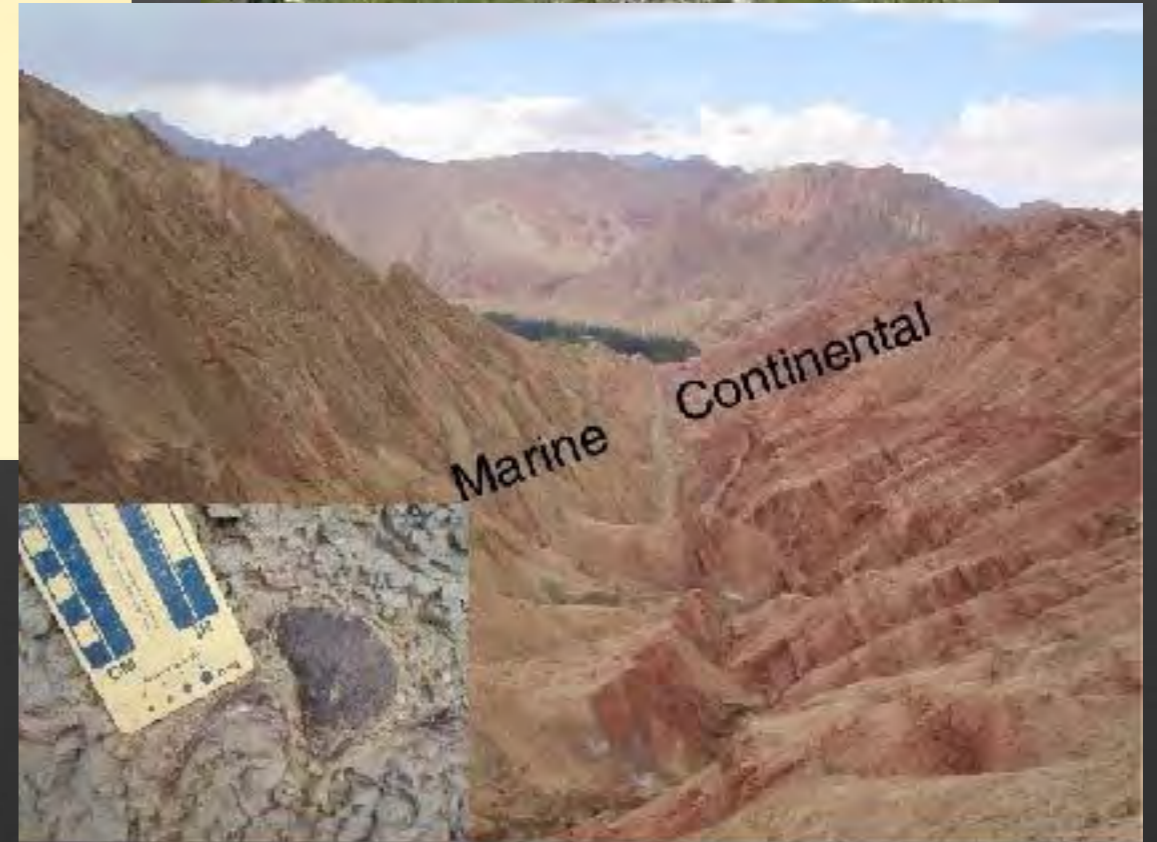
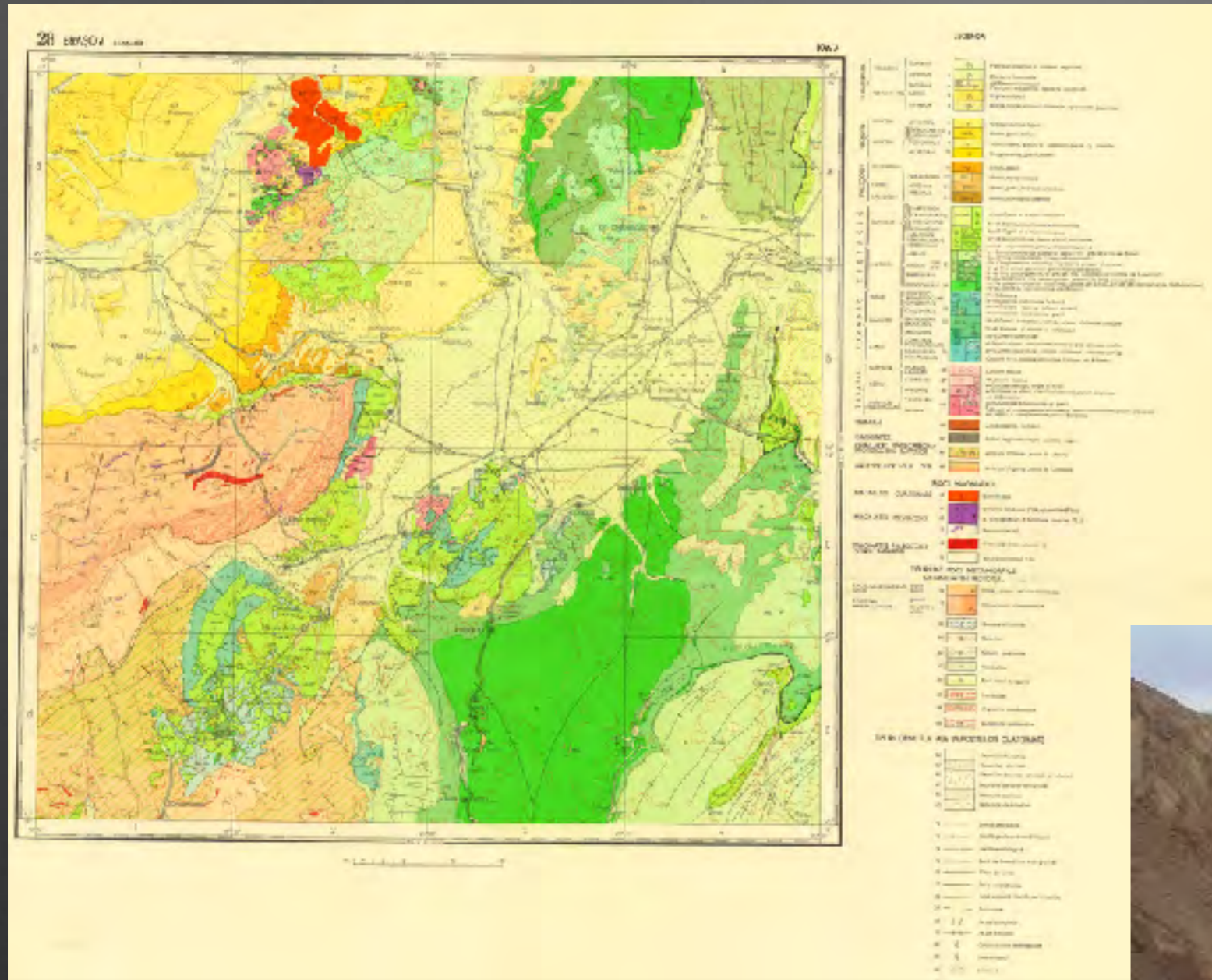
- Early formation of the island arcs (800-400 Ma)
- Big Crush and assembly (400-250 Ma)
- The Alpine Oceans and their islands and volcanoes

All buttressed by older, more rigid blocks

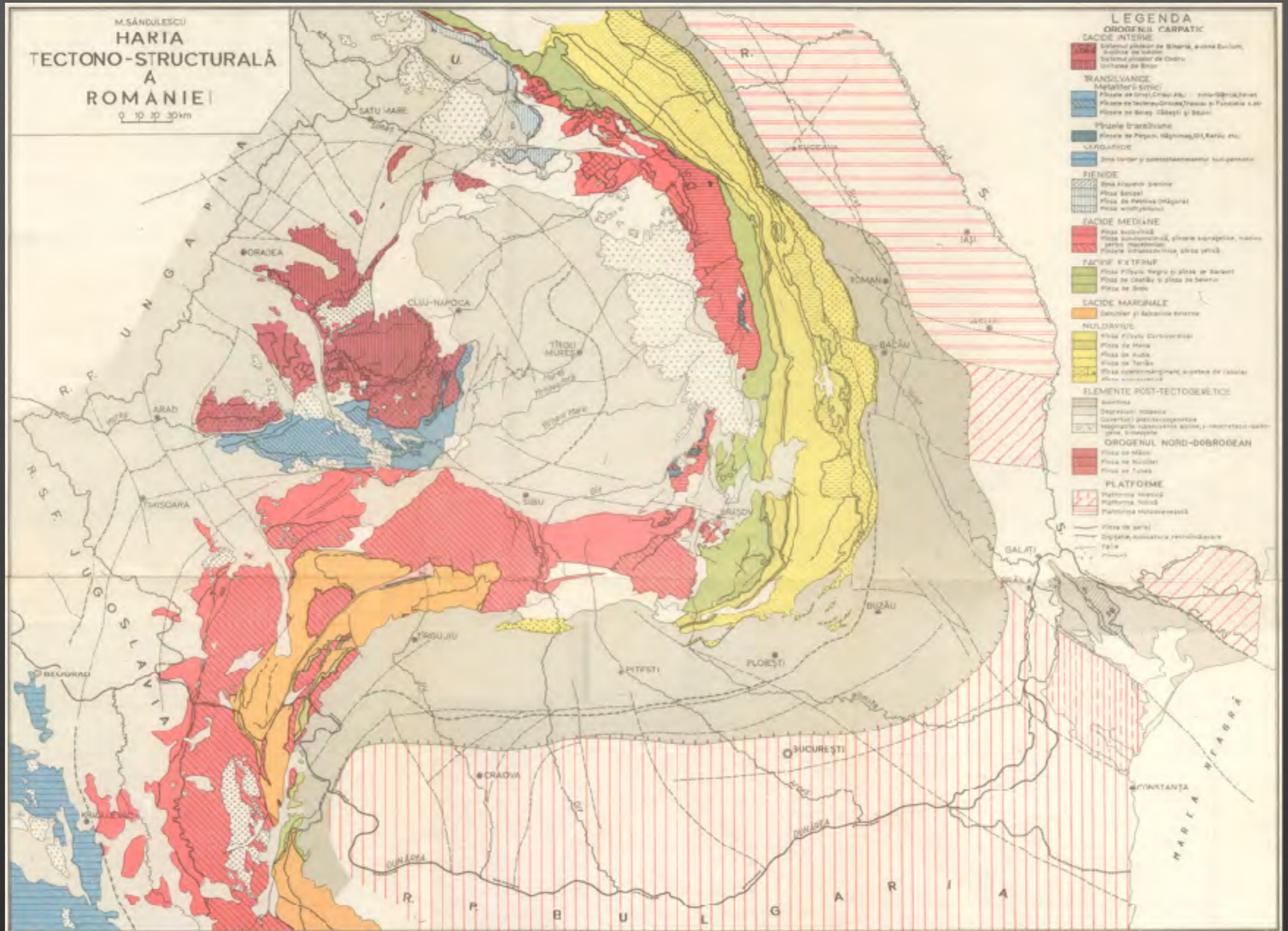
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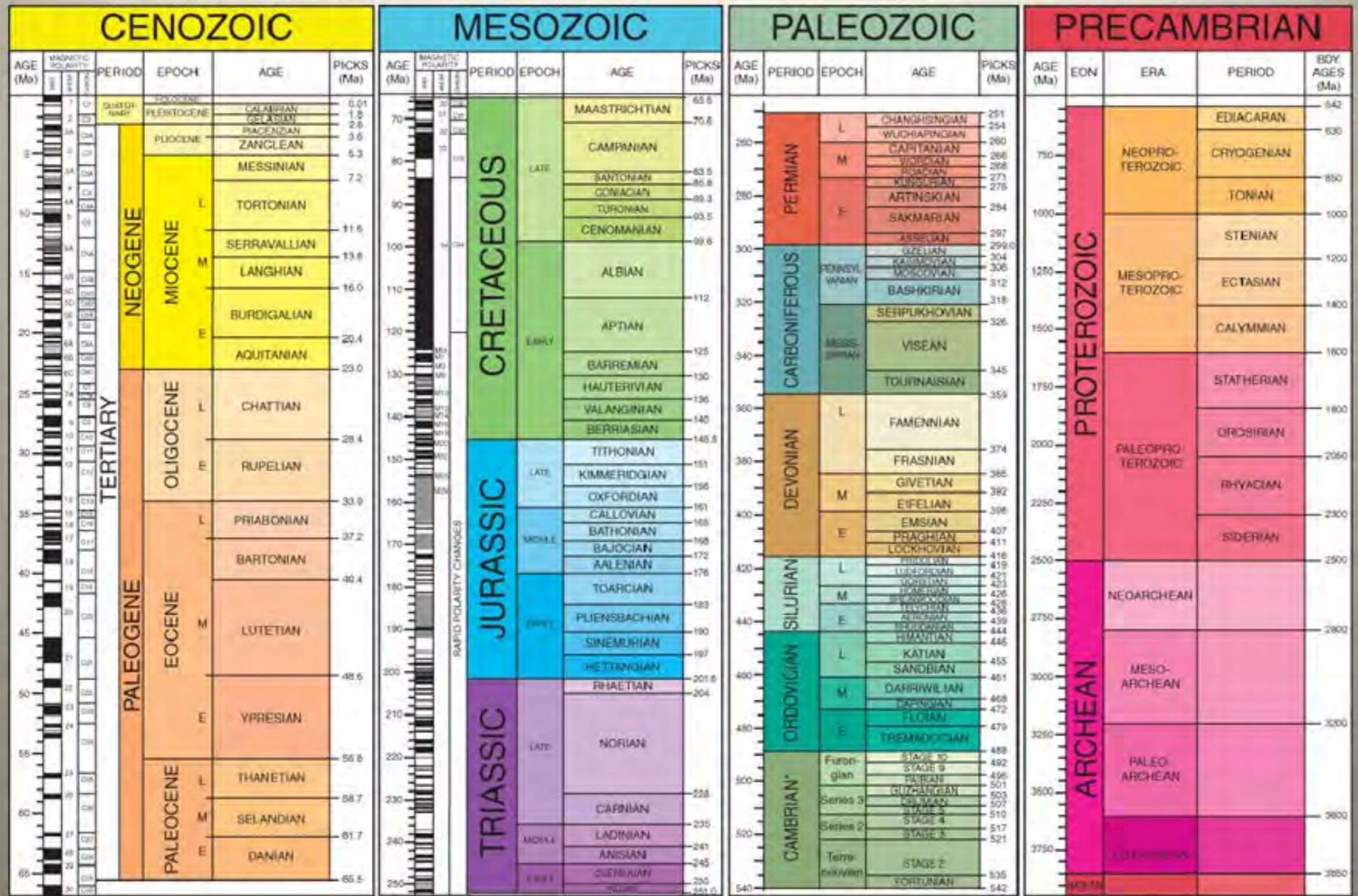








2009 Geologic Time Scale



Subdividing 4.6 billion years into smaller, more “manageable” intervals

Alpine orogeny

Variscan

Caledonian-Cadomian

para-tethys

tethys

big bang

carpy form

TODAY

Phanerozoic

540 Ma

Proterozoic

Archean

2500 Ma

Jack Hills

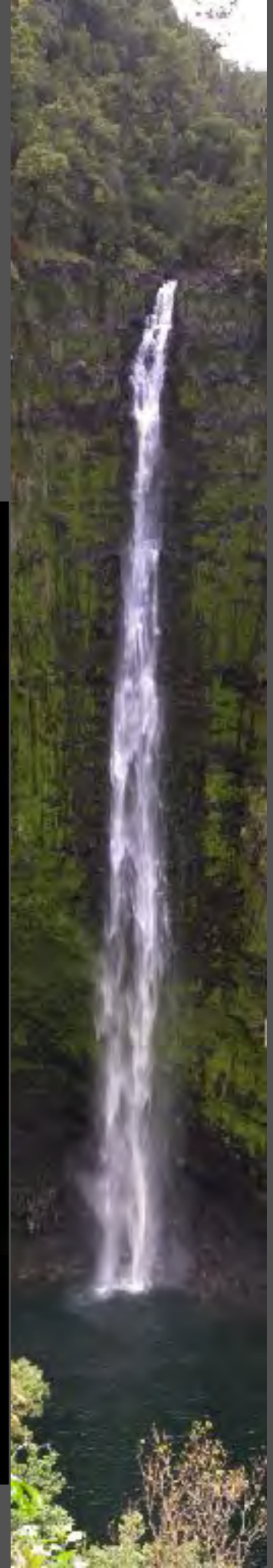
Acasta

Hadean

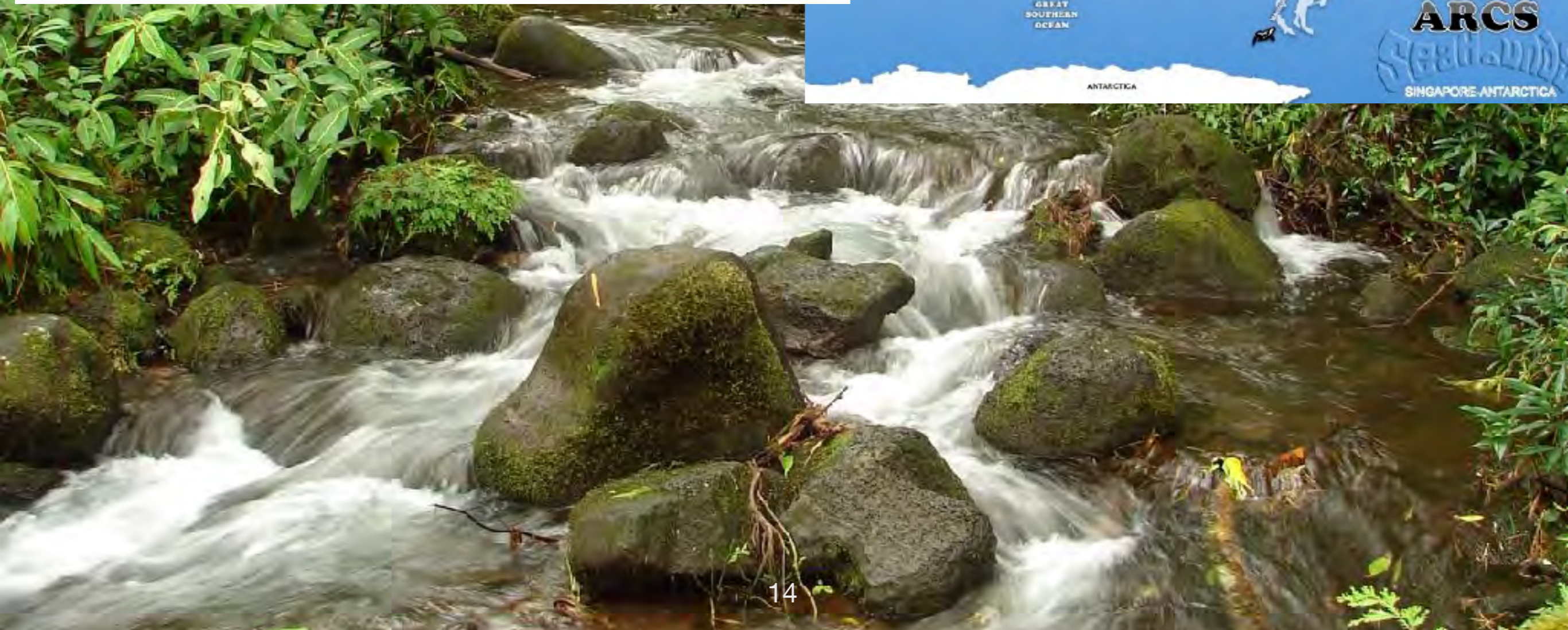
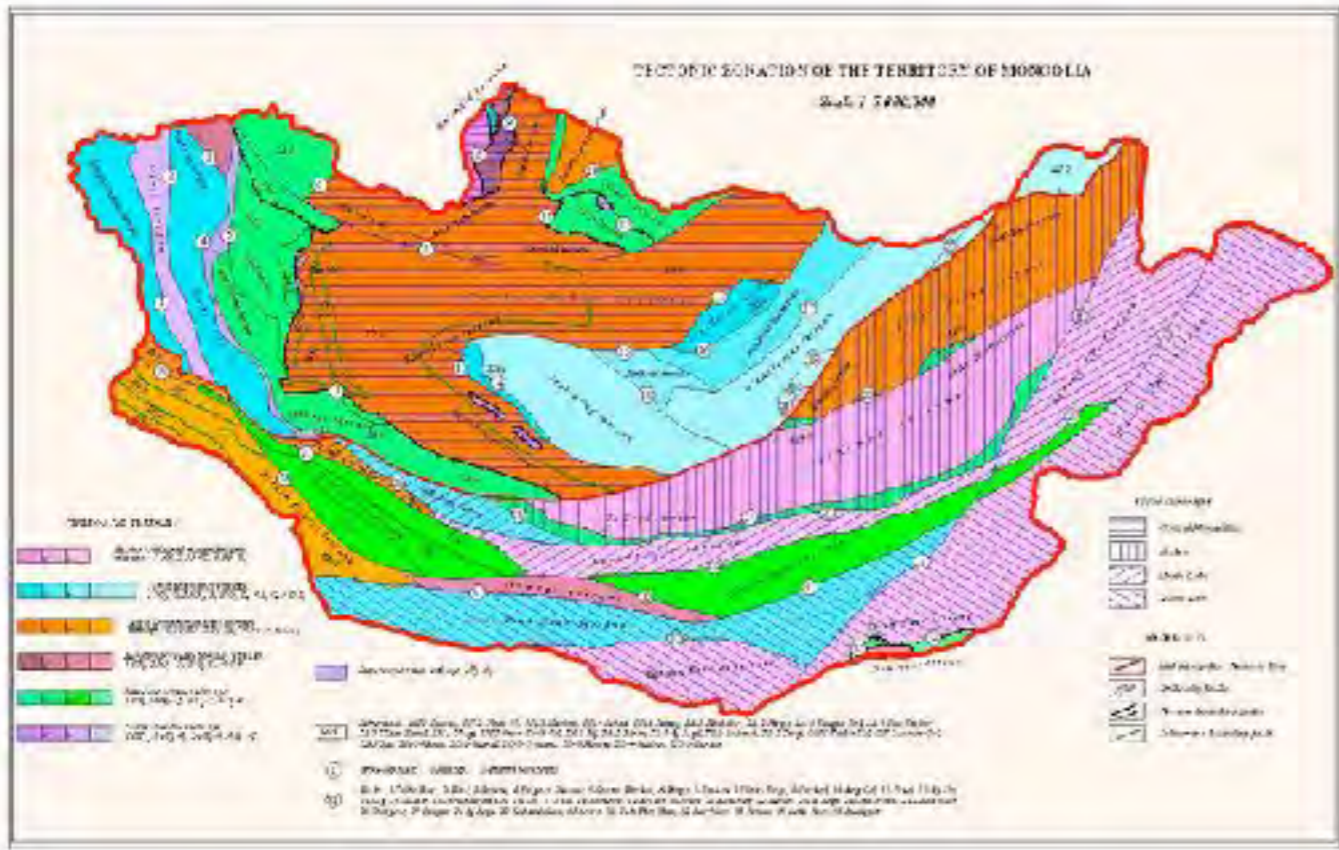
4567 Ma

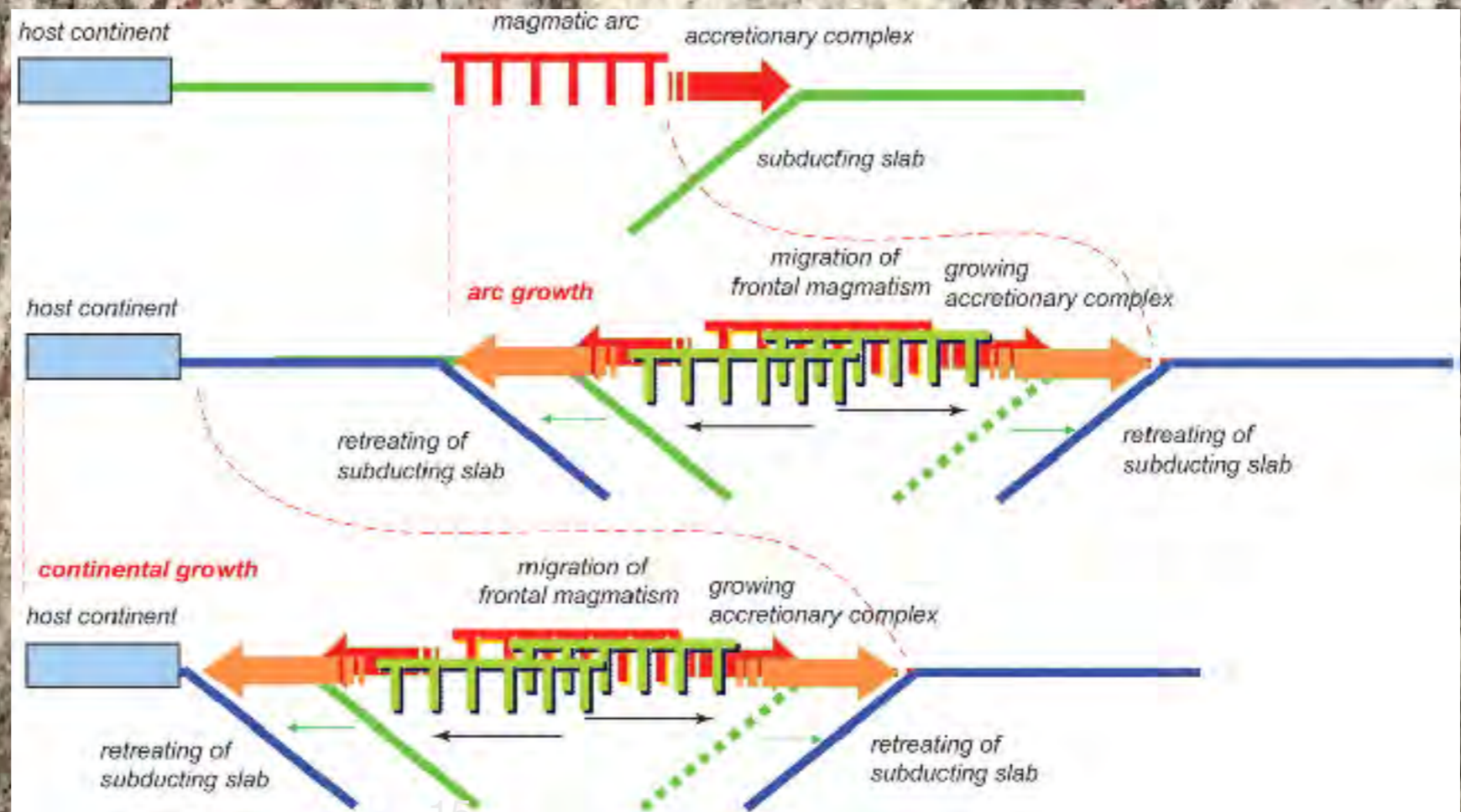
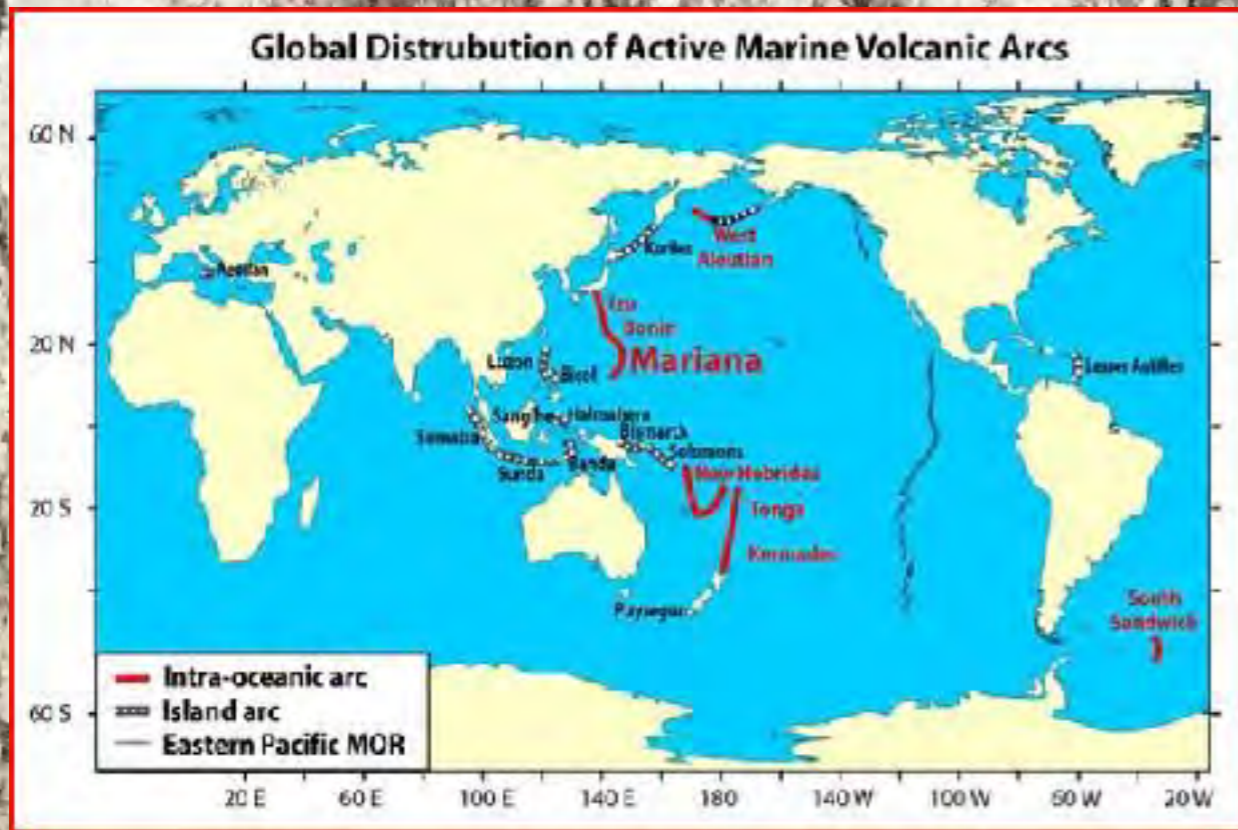
Nuvvuagittuq

4000 Ma

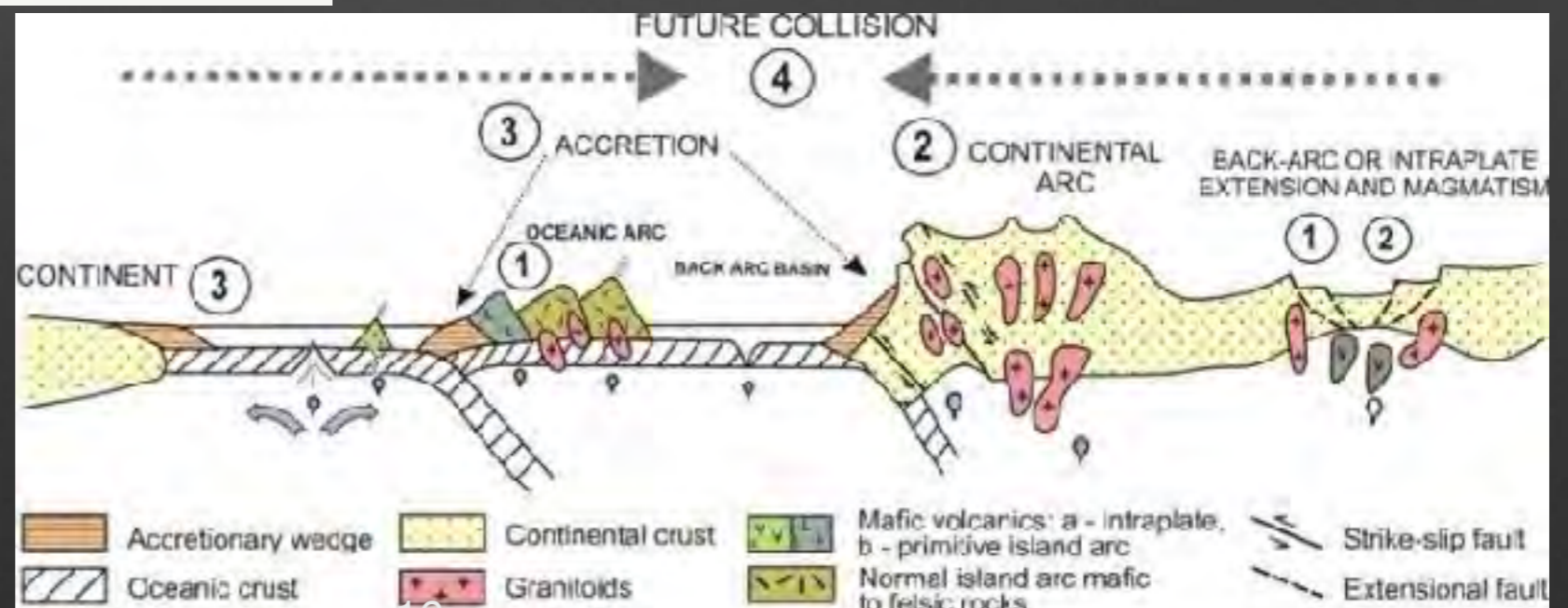
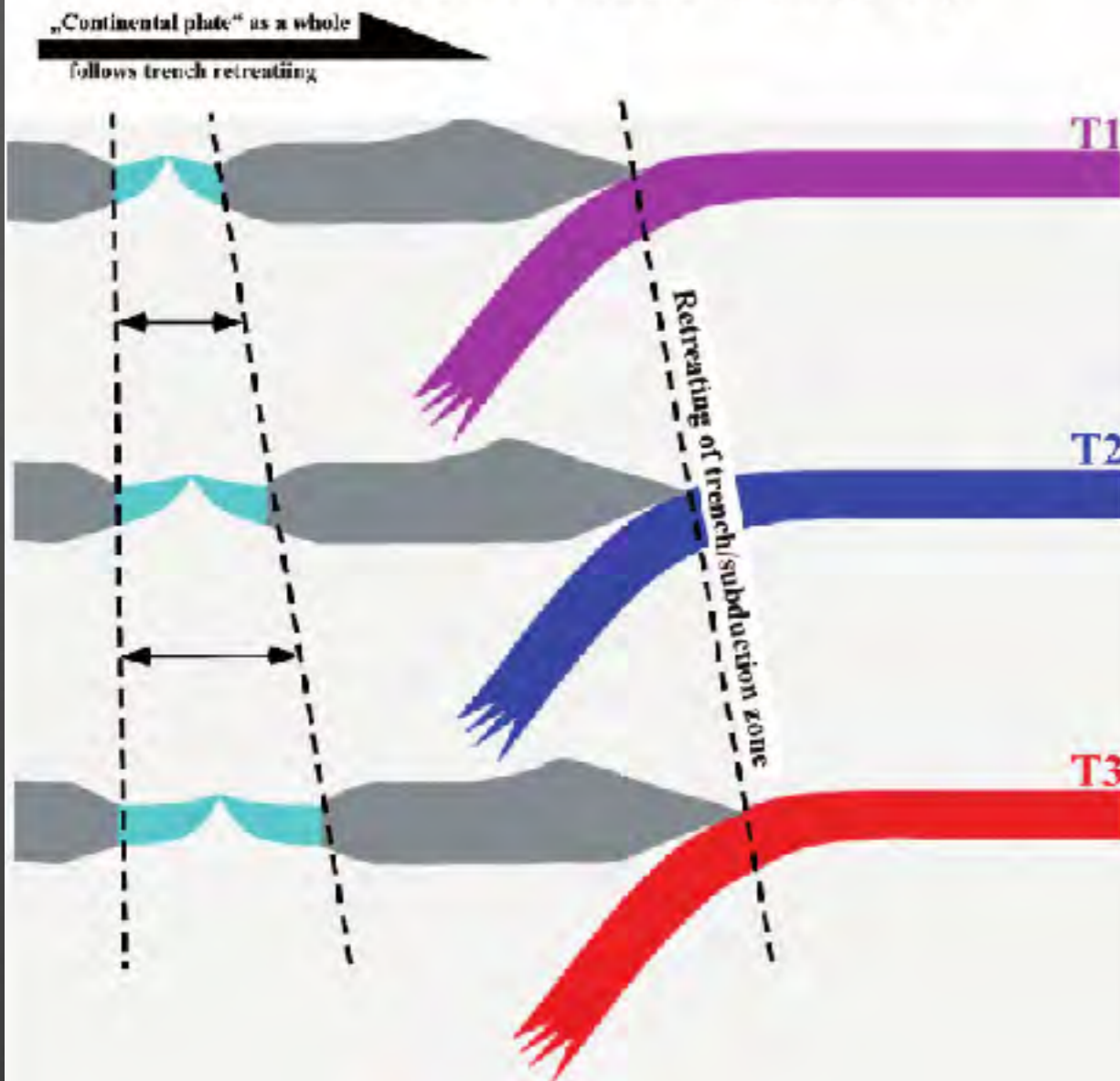


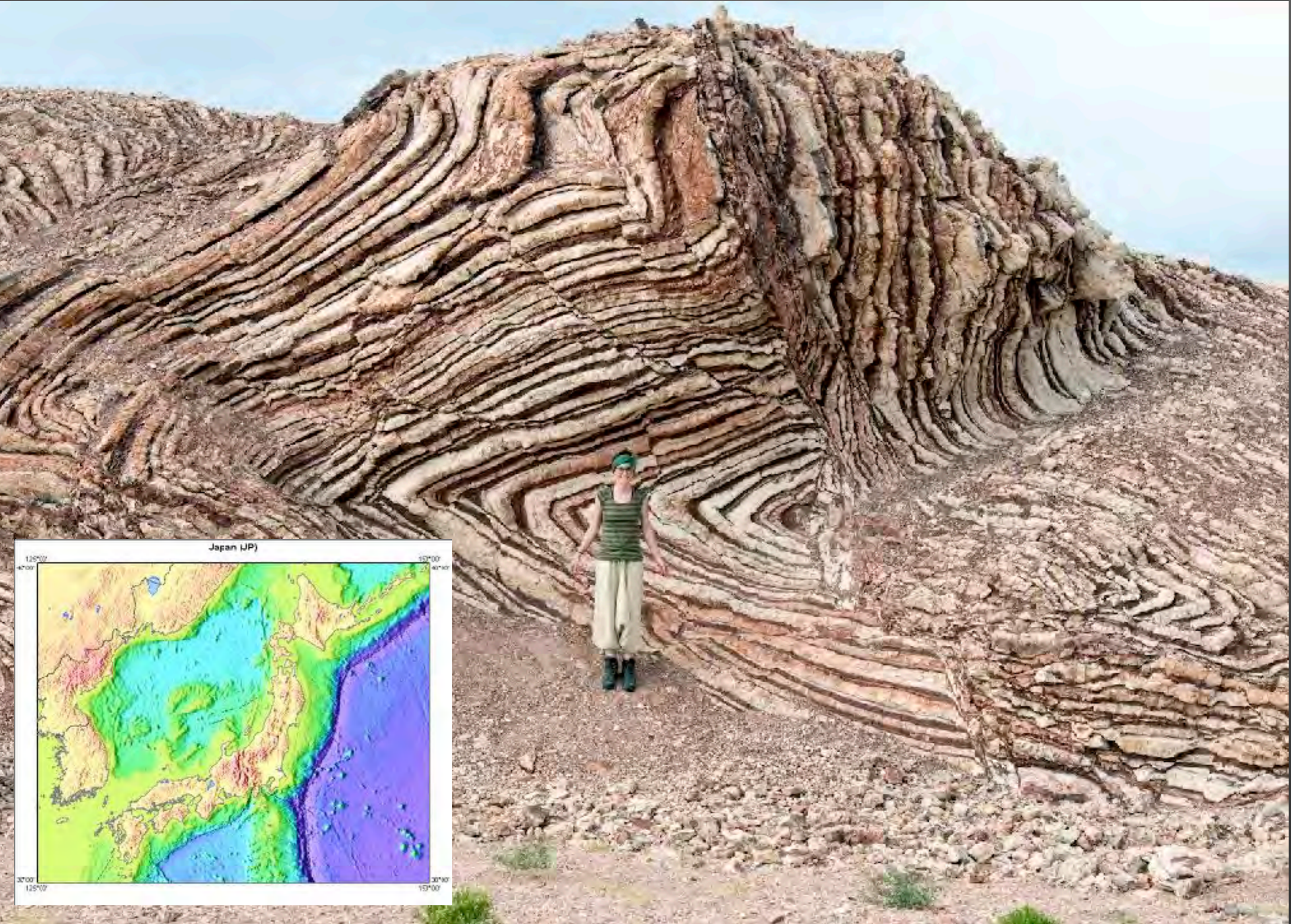






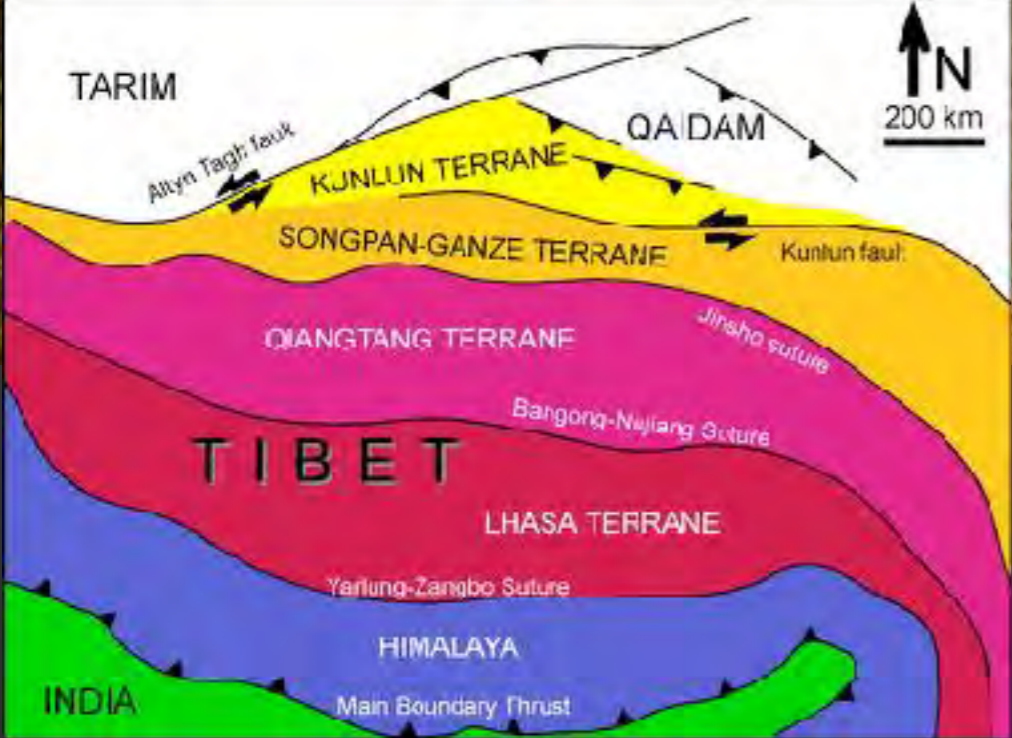
Formation of Back-arc Basin and Spreading



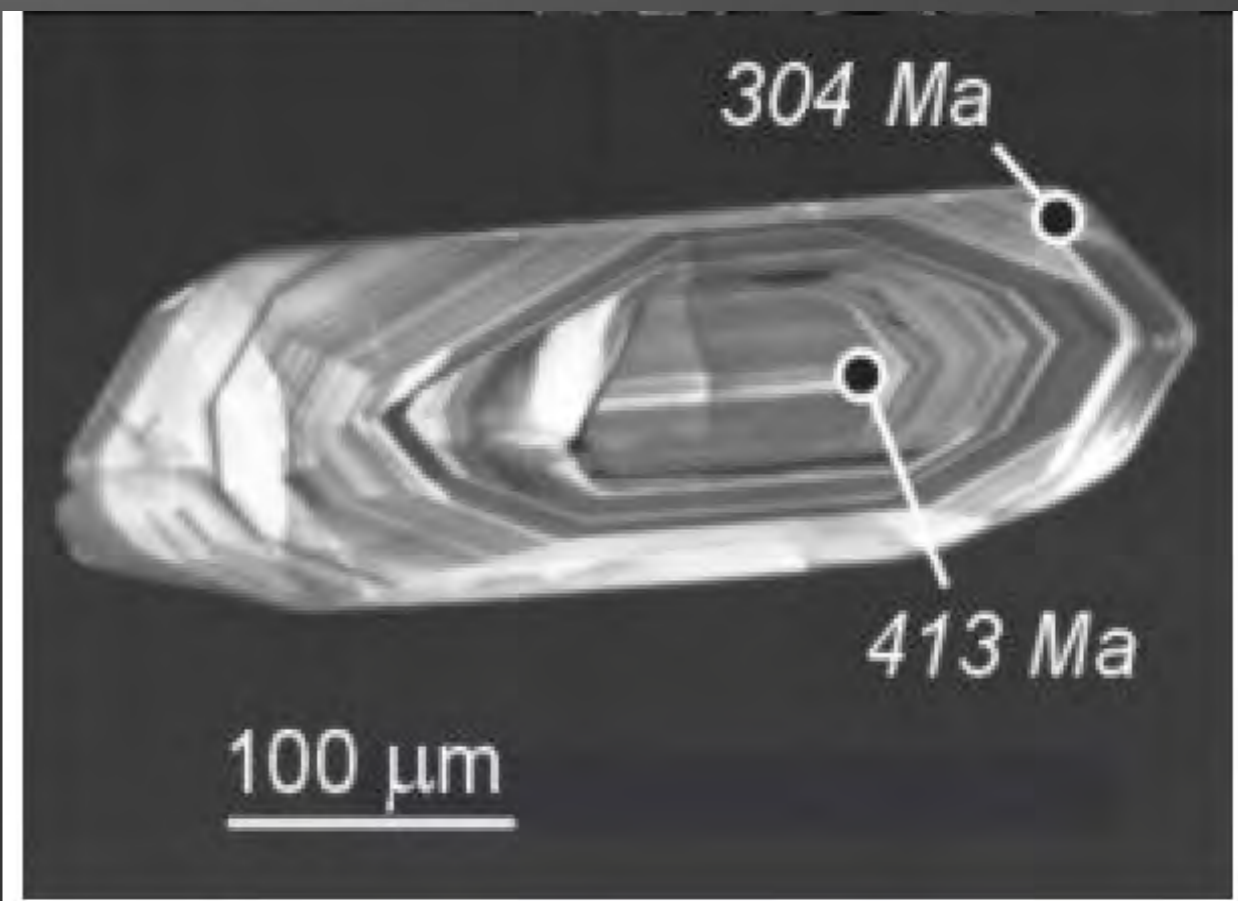
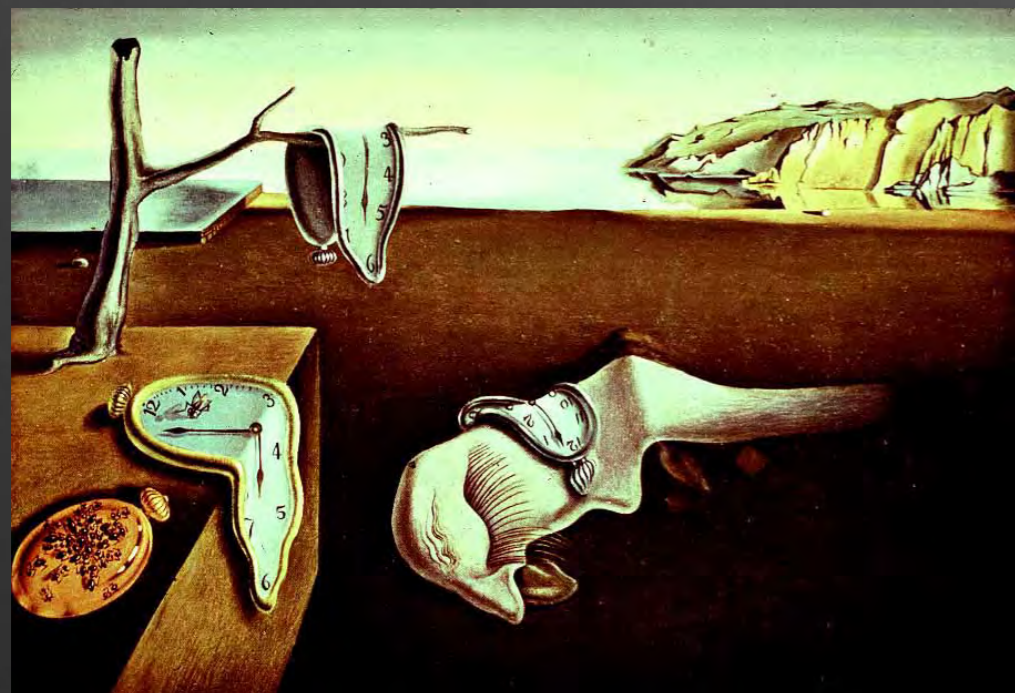
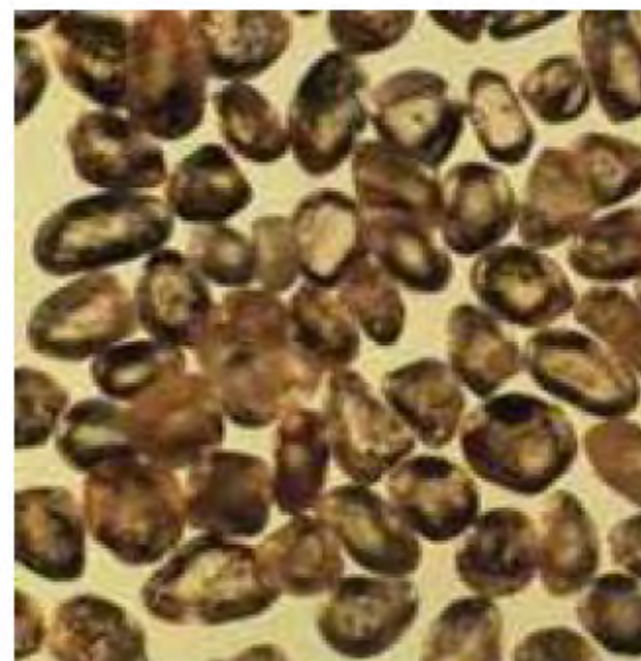
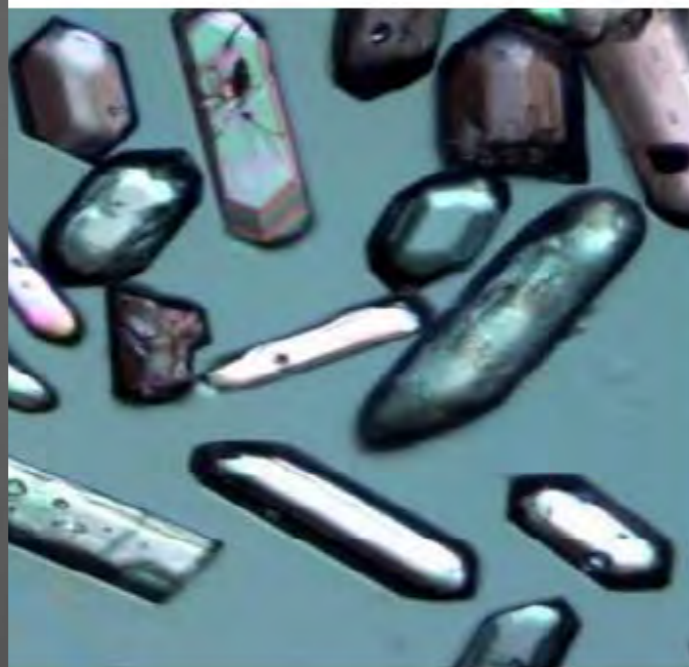
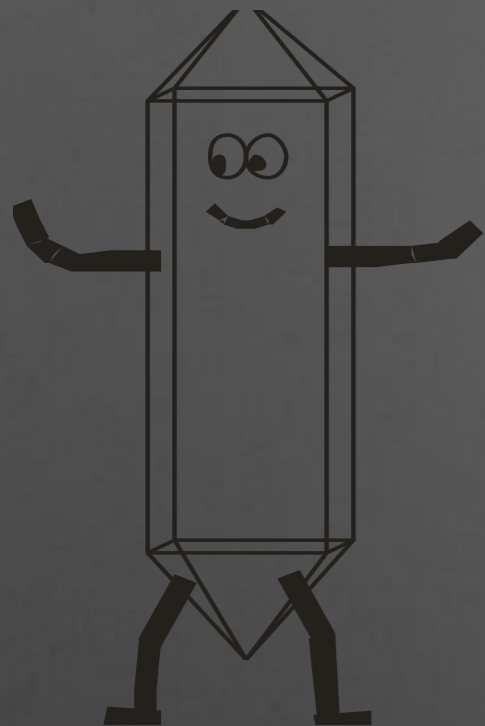


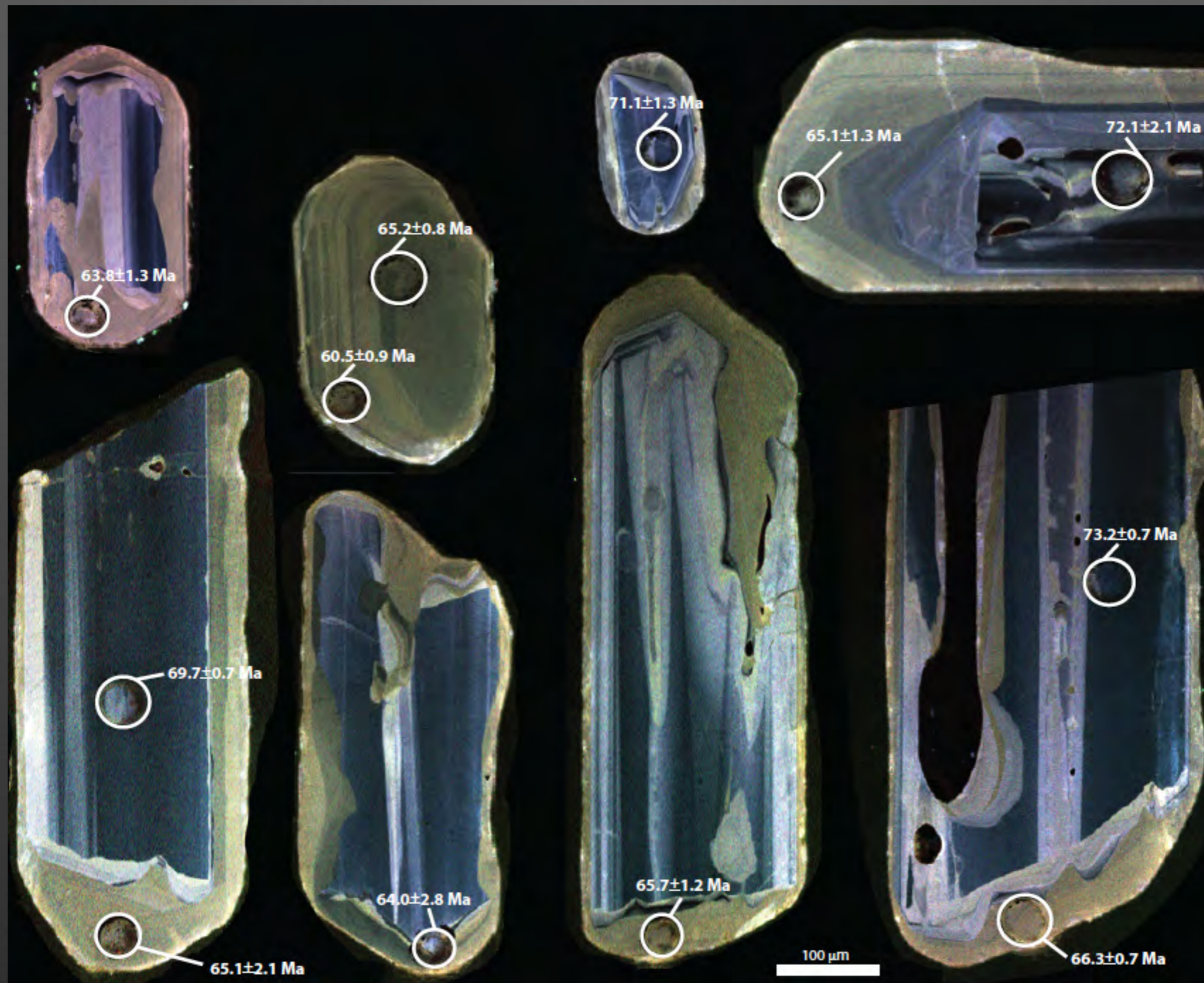


Terrane collage



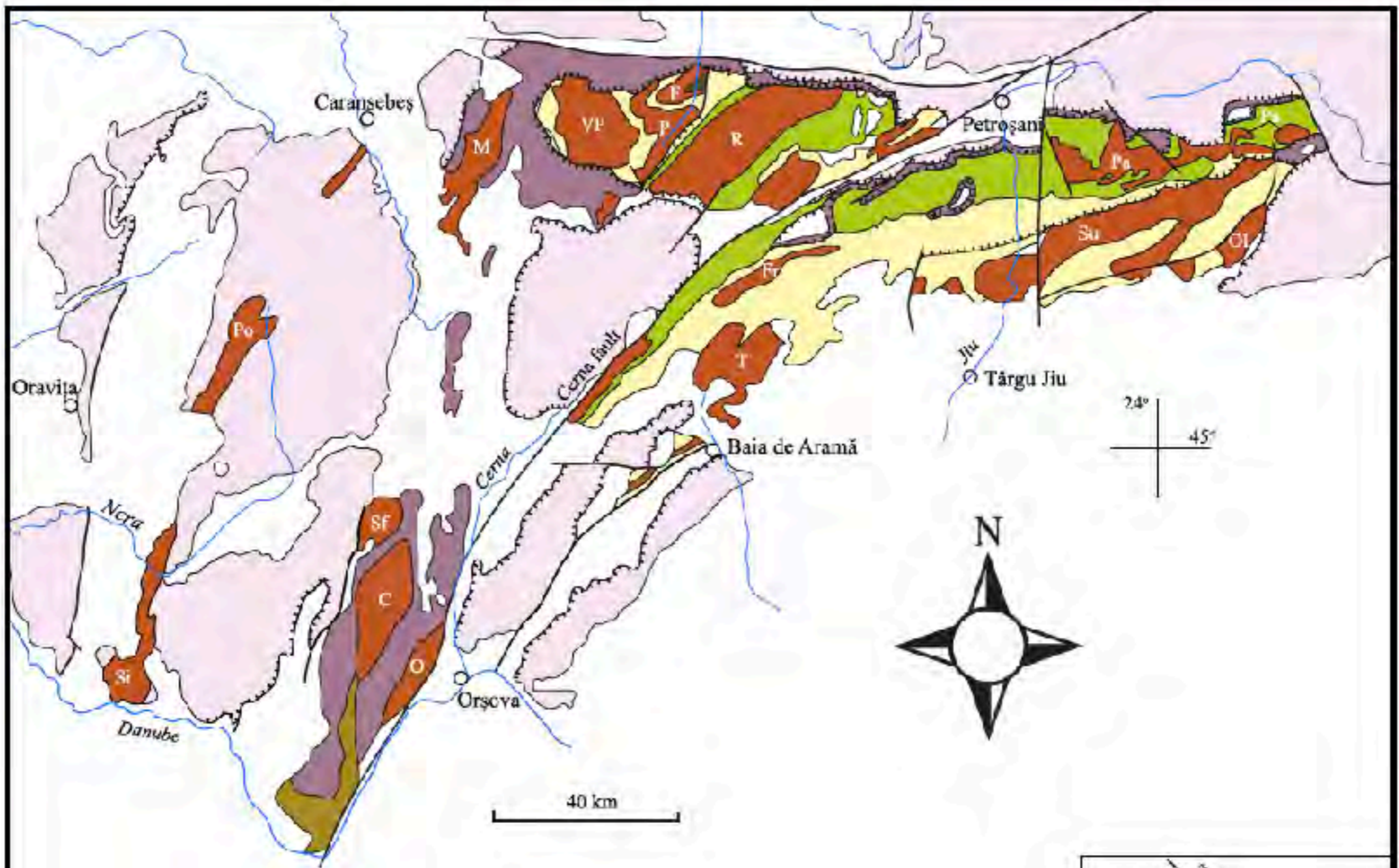






Carpathian basement ages

- Lower Danubian is the oldest - 800 Ma
- Everything else falls into the range of 560-420 igneous ages with inheritance
- 85% of the basement is Cambro Ordovician
- some late Variscan granitoids are yet to be resolved genetically (e.g. Retezat)
- most but not all basement appears metamorphosed post Ordovician



Lower Danubian basement metamorphics

- Drăgșan series
- Lainici-Păuș series
- Rof series

Upper Danubian basement metamorphics

- Tișovița Devonian Ophiolites

Getic-supragetic basement metamorphics

- Granite plutons

Alpine thrust

Fault

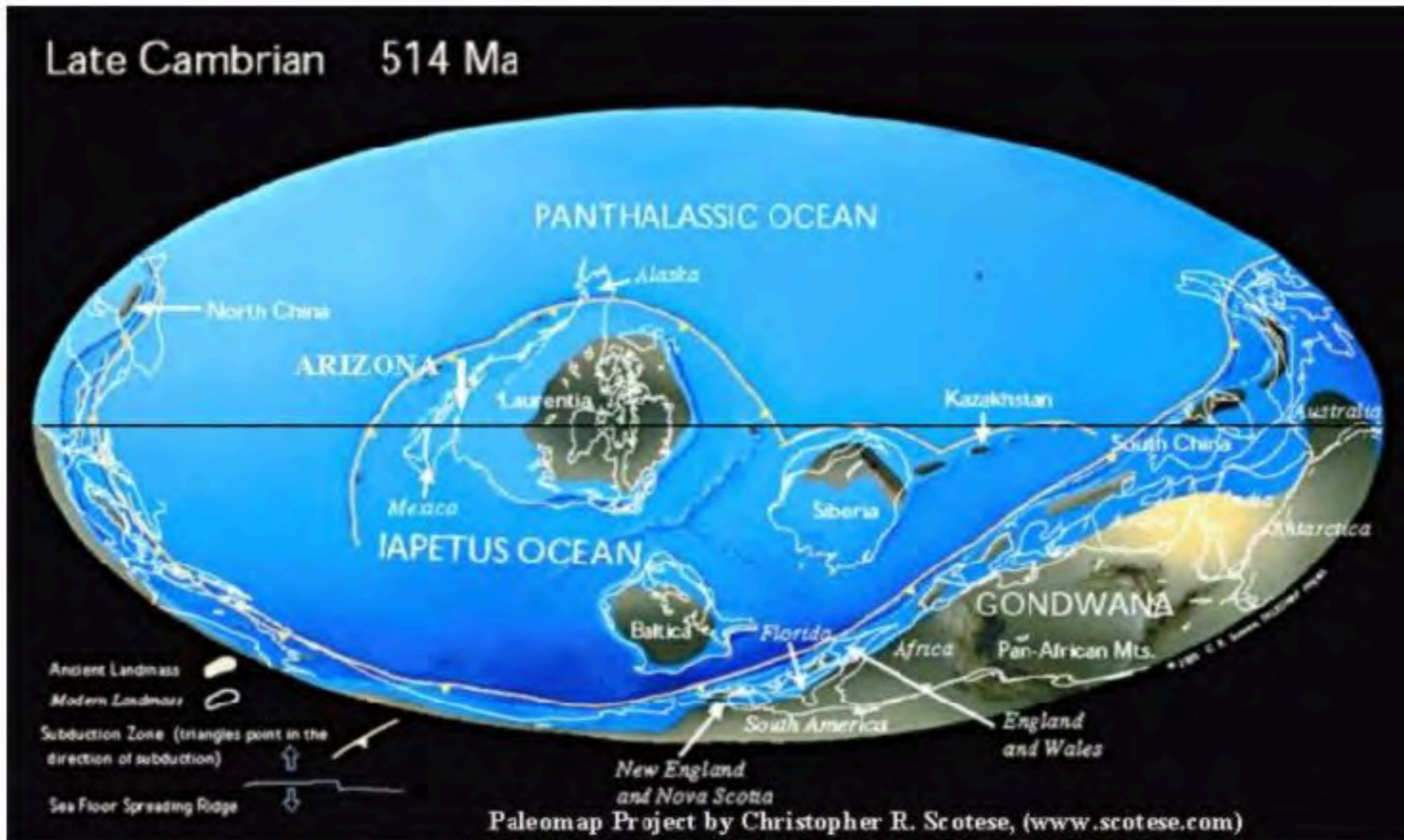




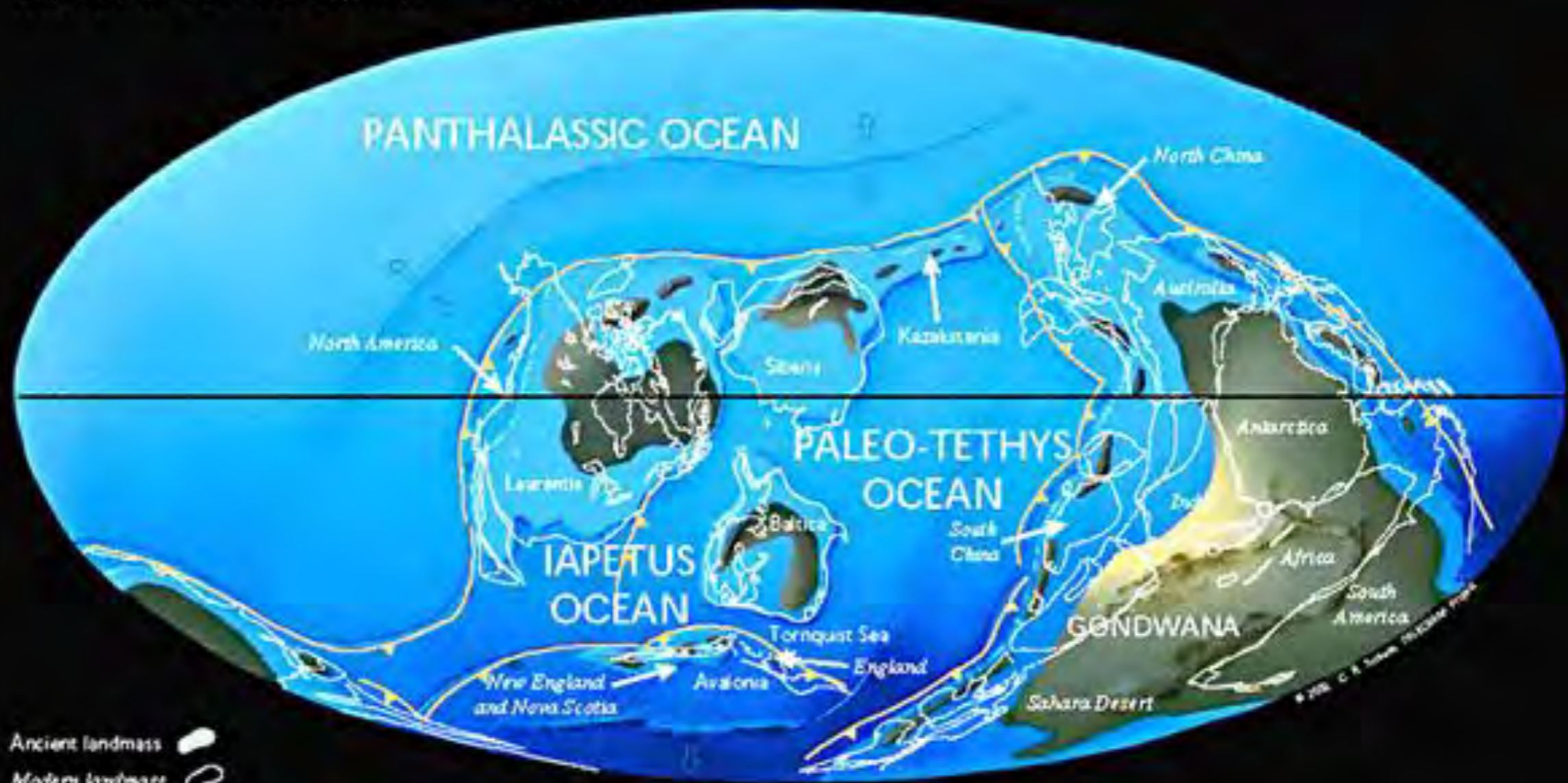




Late Cambrian 514 Ma



Middle Ordovician 458 Ma

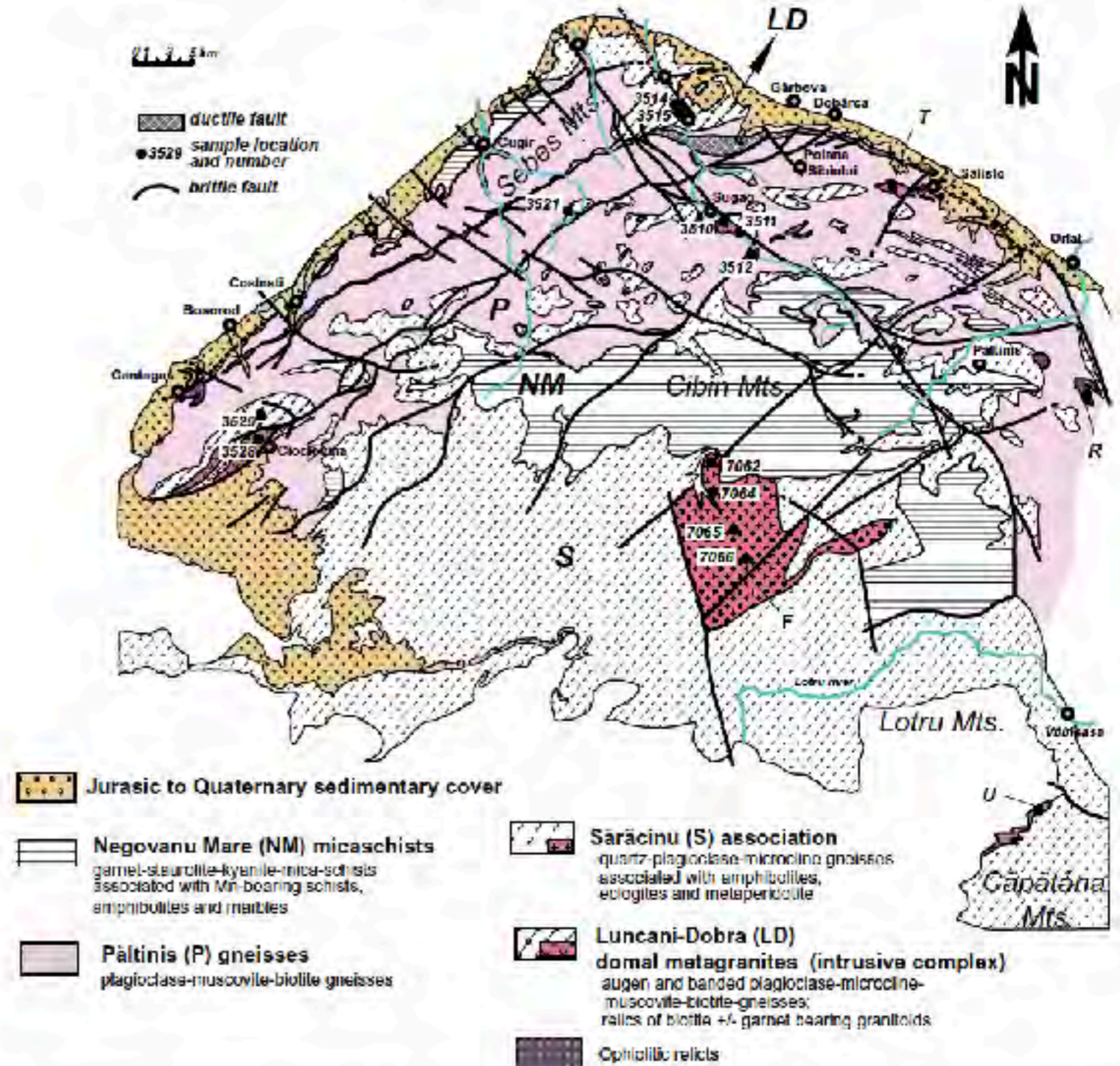


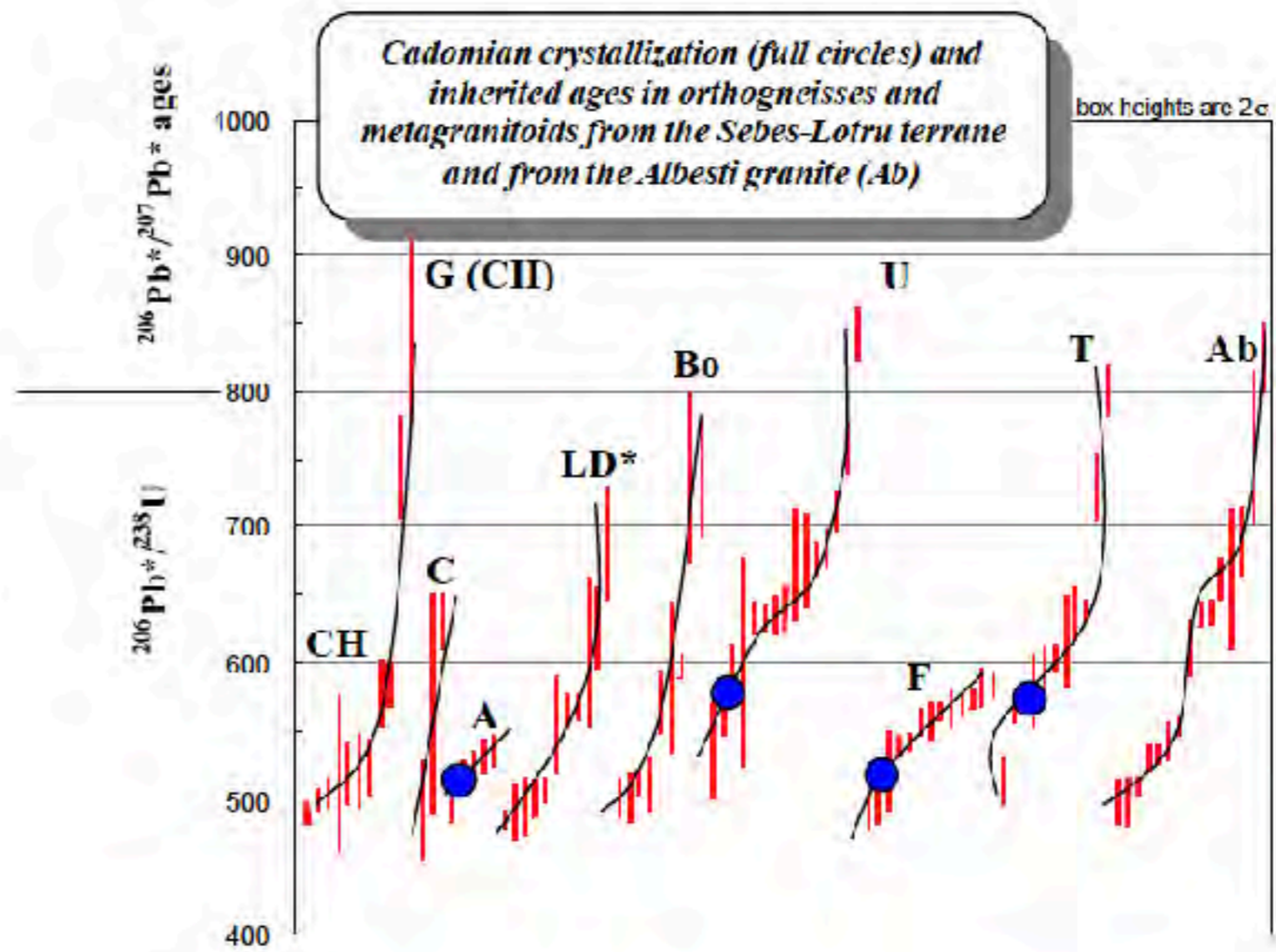
- Ancient landmass
- Modern landmass
- Subduction Zone (triangles point in the direction of subduction)
- Sea Floor Spreading Ridge

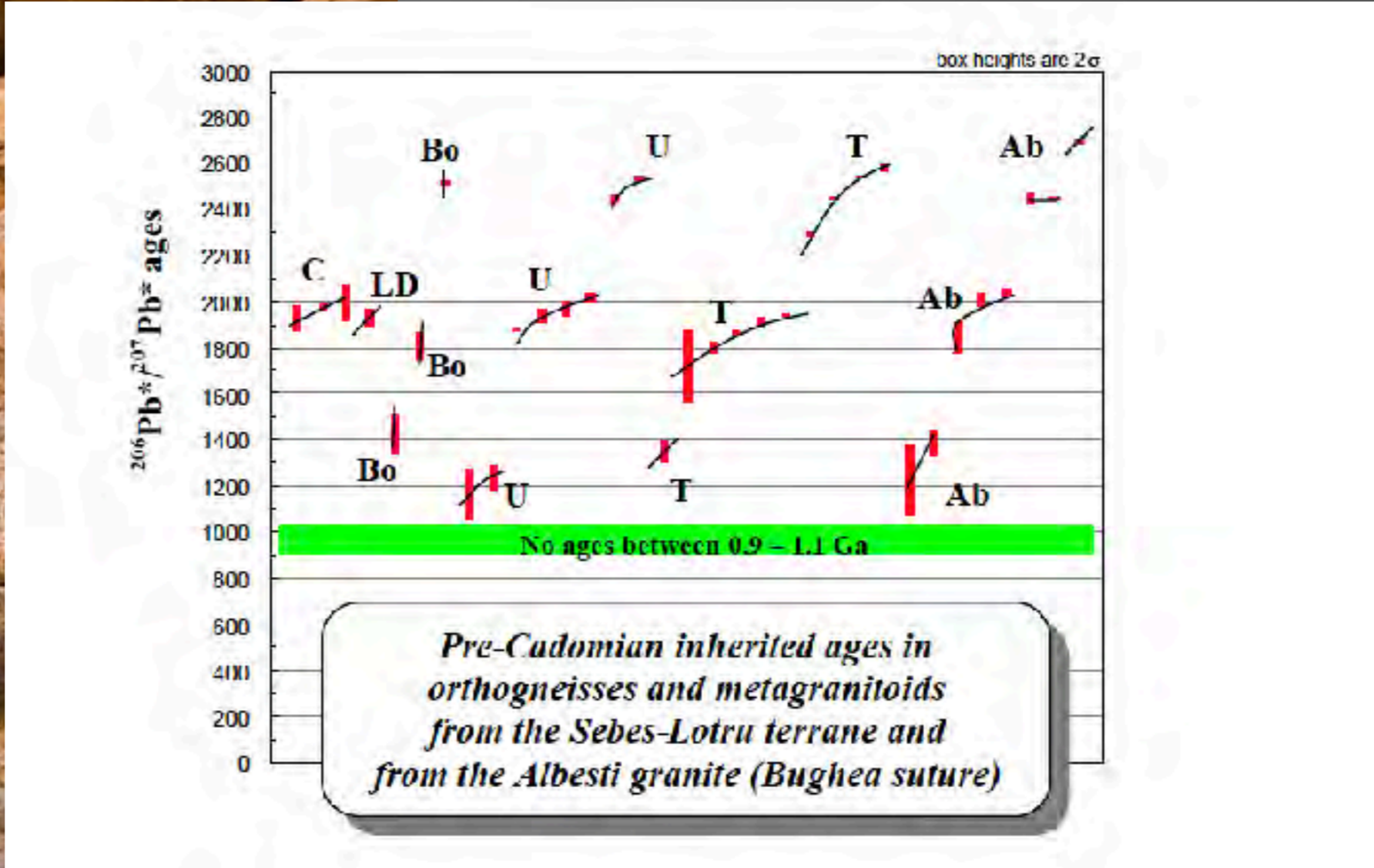


Geological map of the Sebes-Cibin Mountains with main lithology of the Sebes-Lotru terrane and samples location

(after Stelea, I., 1999, modified according to Sabau, 1994, Sabau and Massone, 2003 and Geological Map of Romania scale 1:200 000, Orastie sheet)



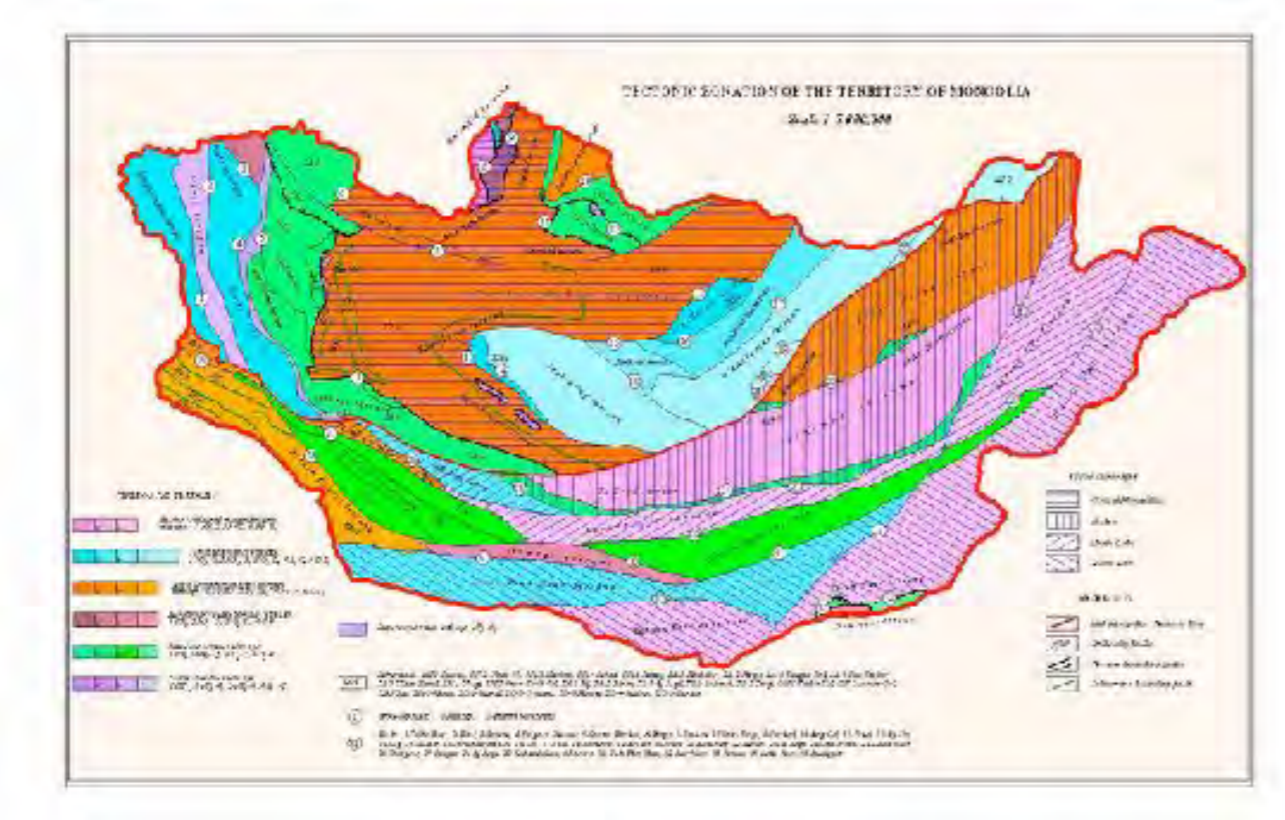
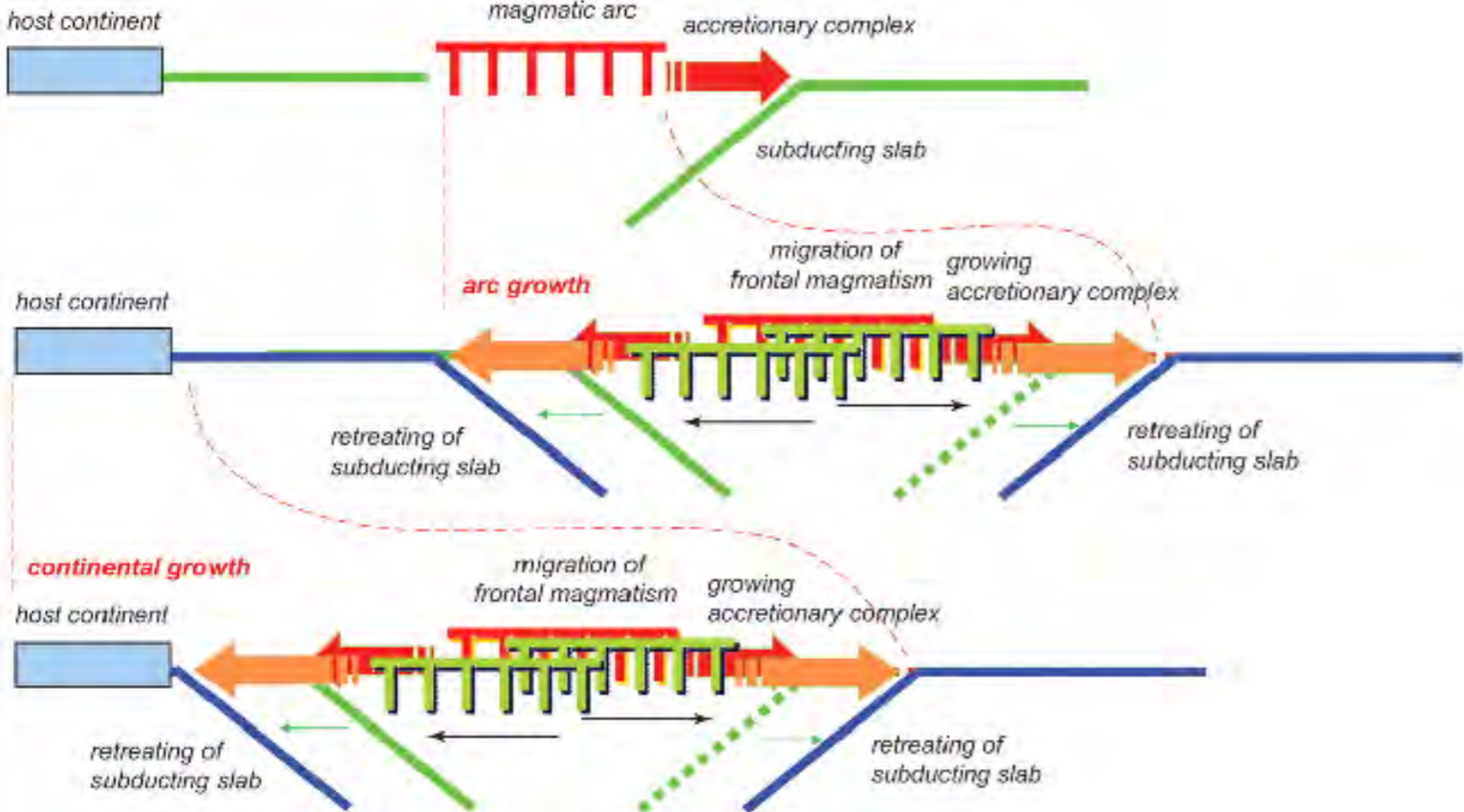




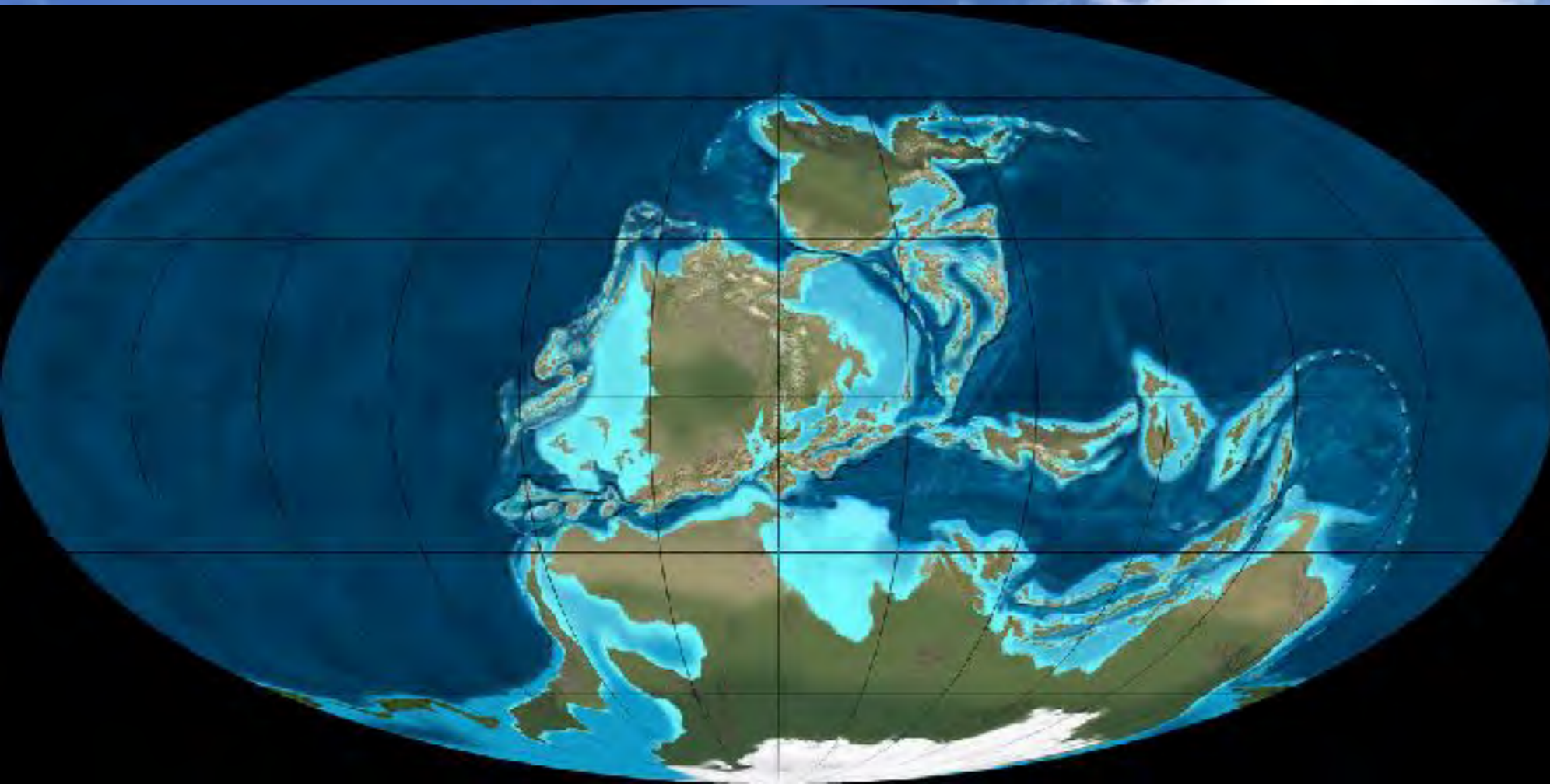


Age of the big bang?

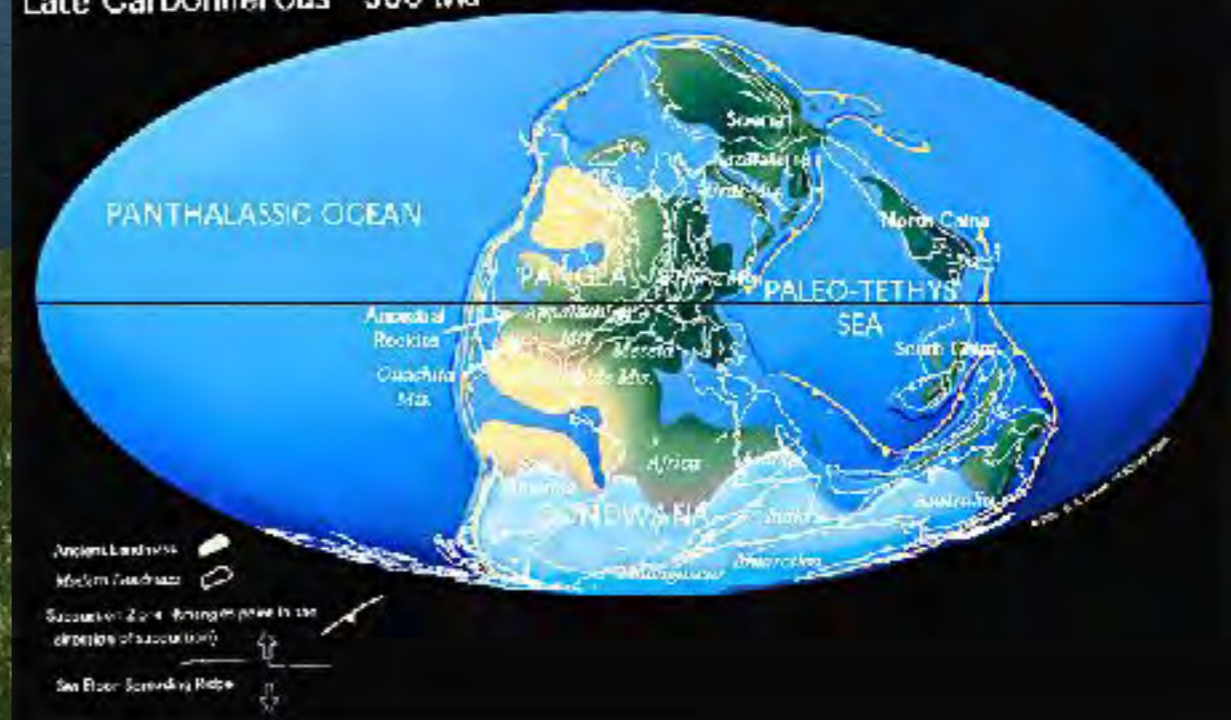
- Not all metamorphic terrains have the evidence for this hard collision
- The high grade terrains do; the biggest of them is the Sebes Lotru unit in the South Carpathians;
- Think Fagaras metamorphic rocks
- and think Himalayan style collision
- It took place between 350-325 Ma and it was followed by extensional collapse

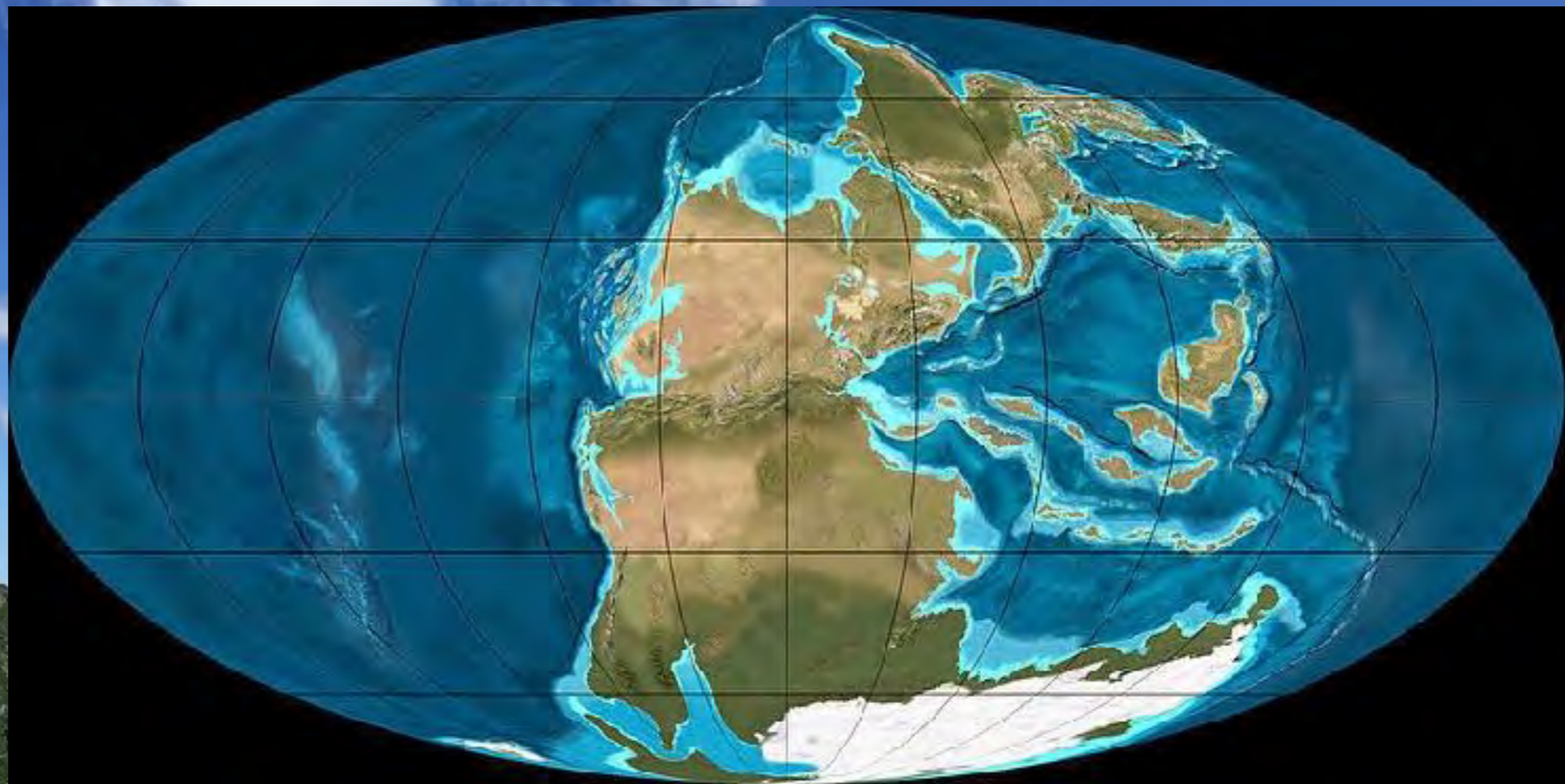




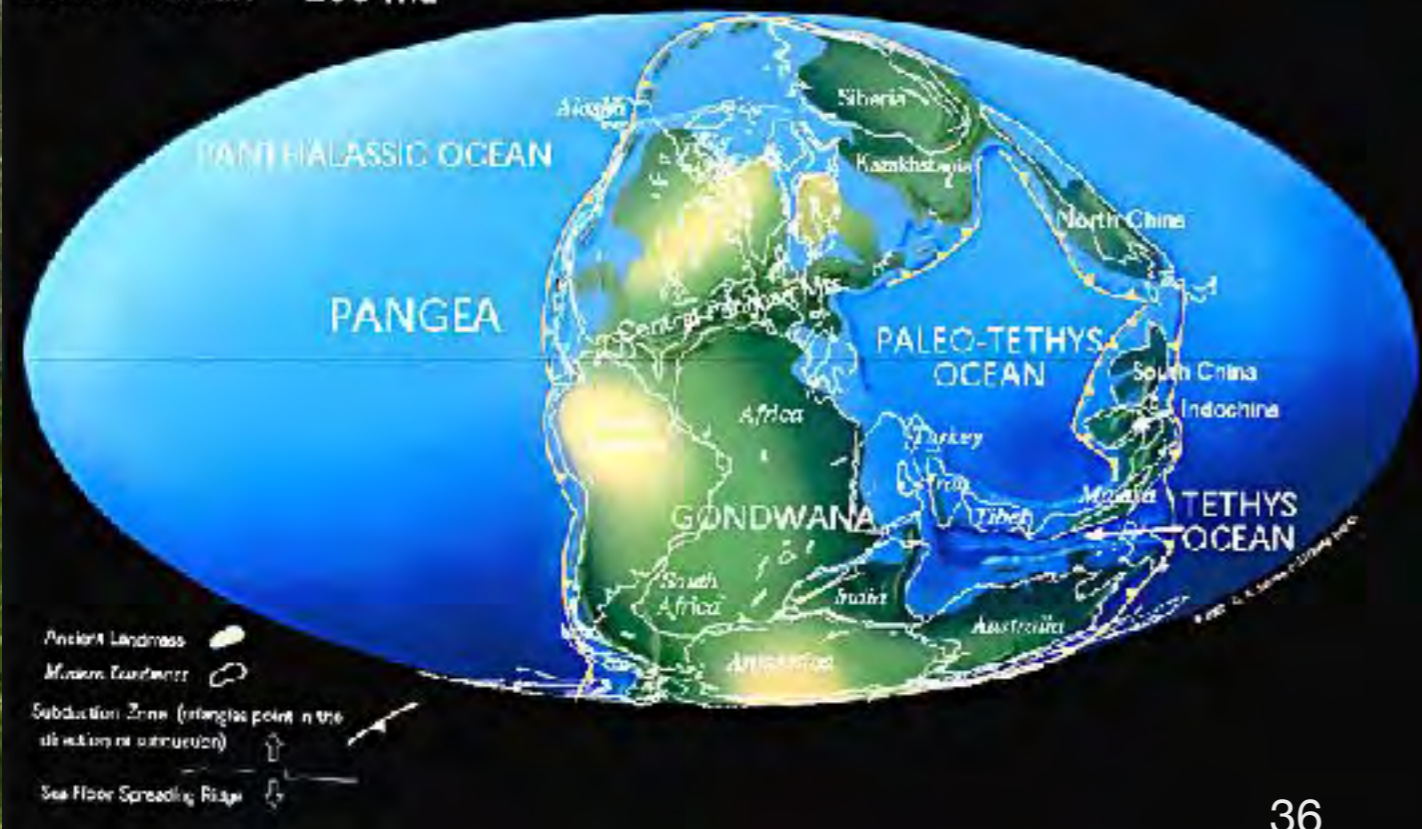


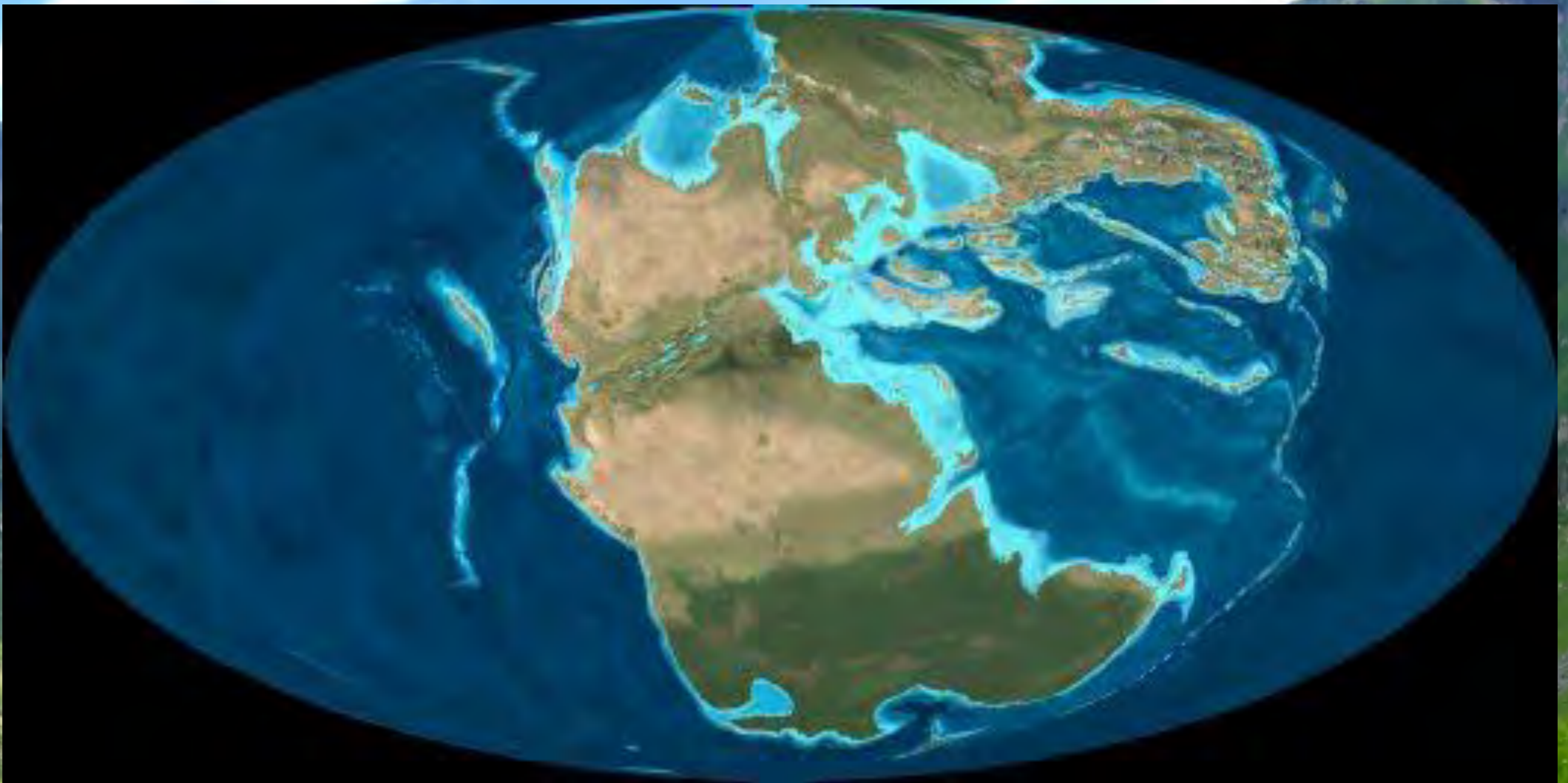
Late Carboniferous 306 Ma



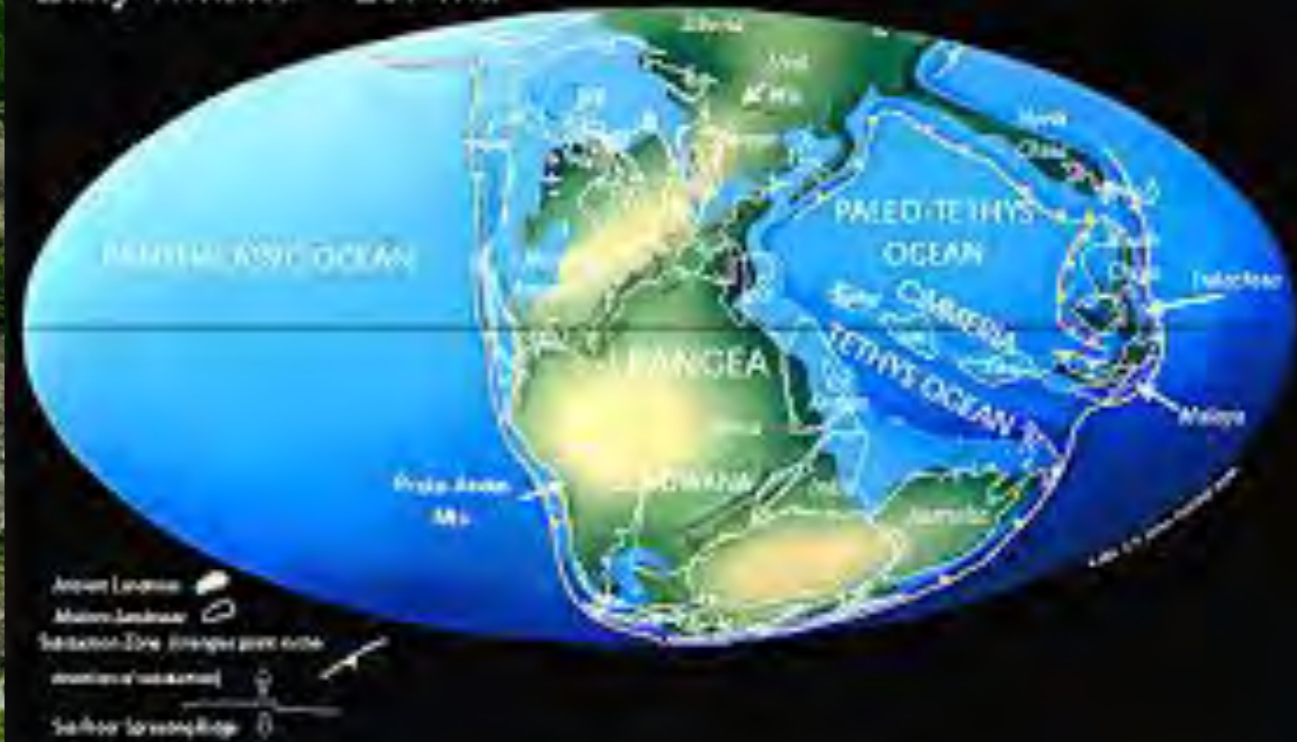


Late Permian 255 Ma



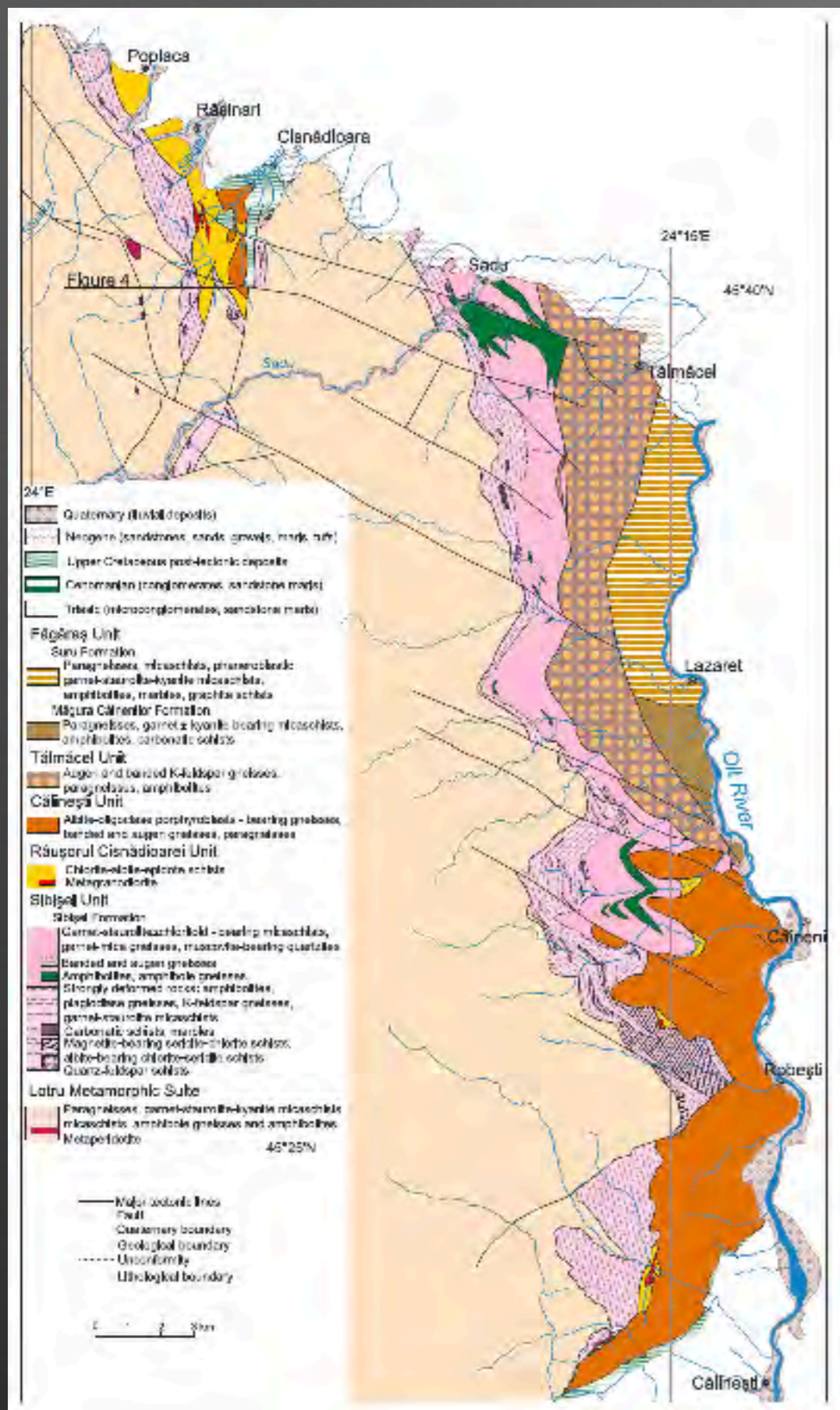


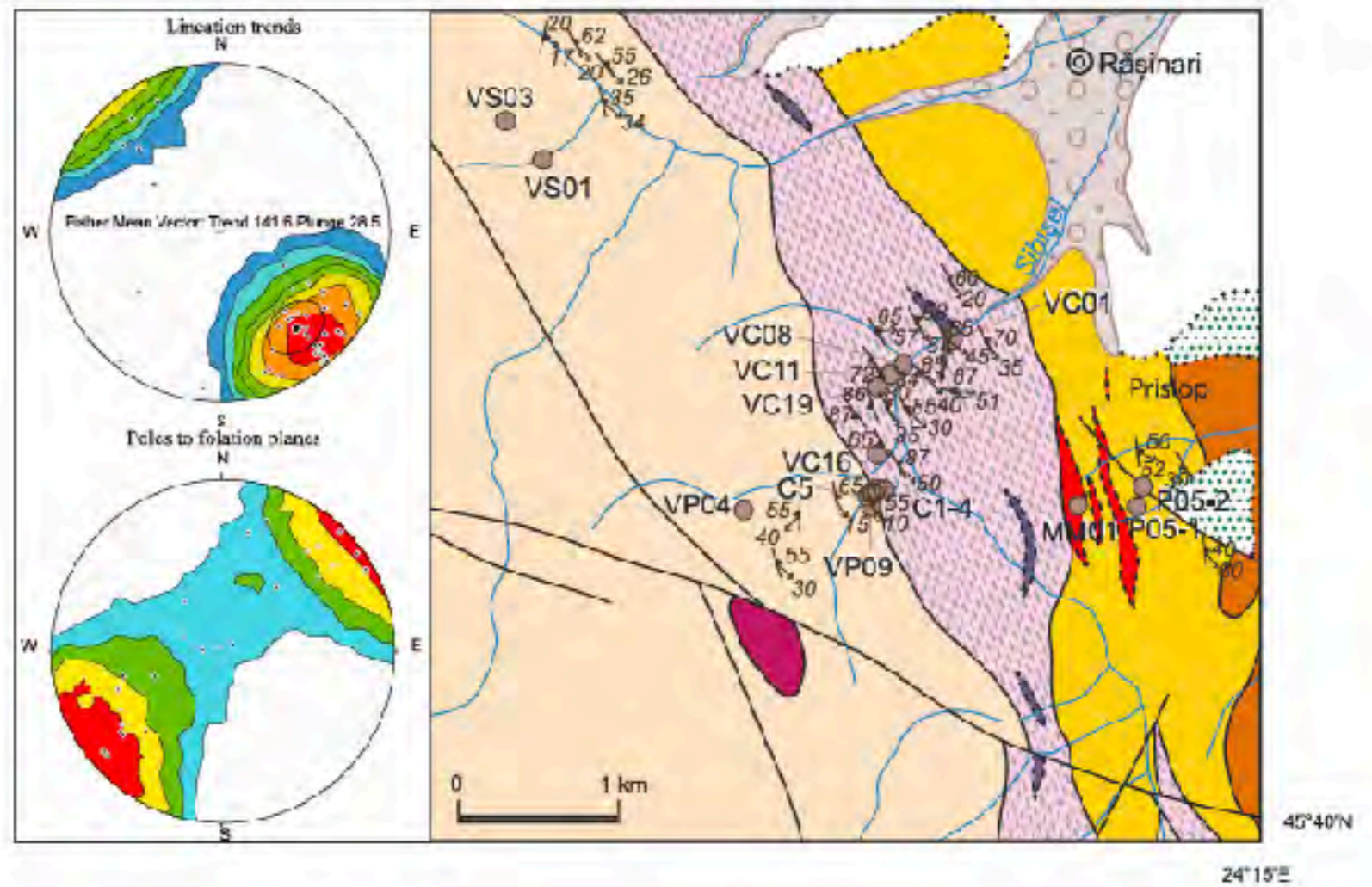
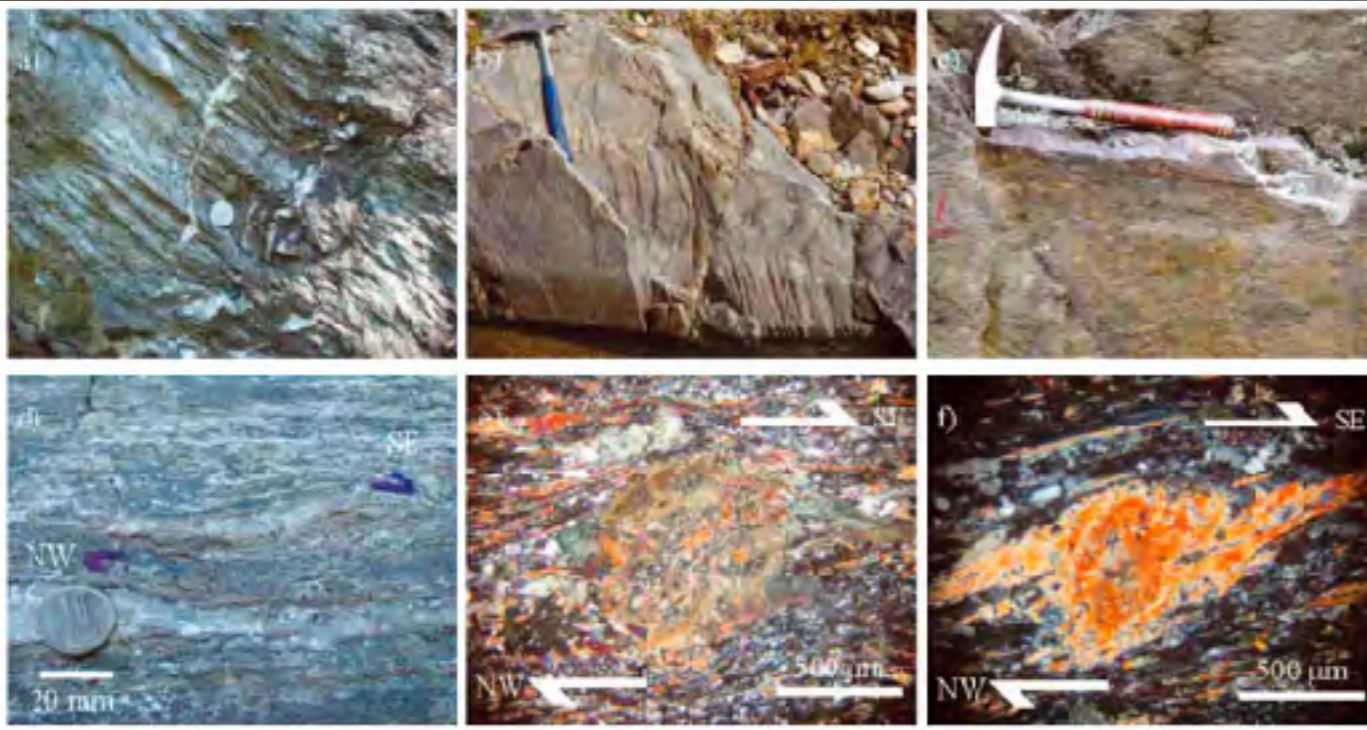
Early Triassic 237 Ma



Example- Rasinari

- Major shear zone
- puts in contact 2 different types of metamorphic assemblages, on high grade and another one low grade
- Both are island arc terranes with Cambro Ordovician ages
- They were attached to each other across an oceanic-like material in the Permian
- Only one (the high grade) had Variscan collisional (big bang) metamorphism
- The other has been floating across the paleo-tethys until its ultimate collision in the Permian, possibly even the early Triassic





Summary

- The basement formed as volcanic arcs and marginal basins outboard of paleo Africa and collided with the east European craton in late Paleozoic as the Paleo Tethys ocean and its variants were closing. Some units underwent HARD others SOFT collisions.

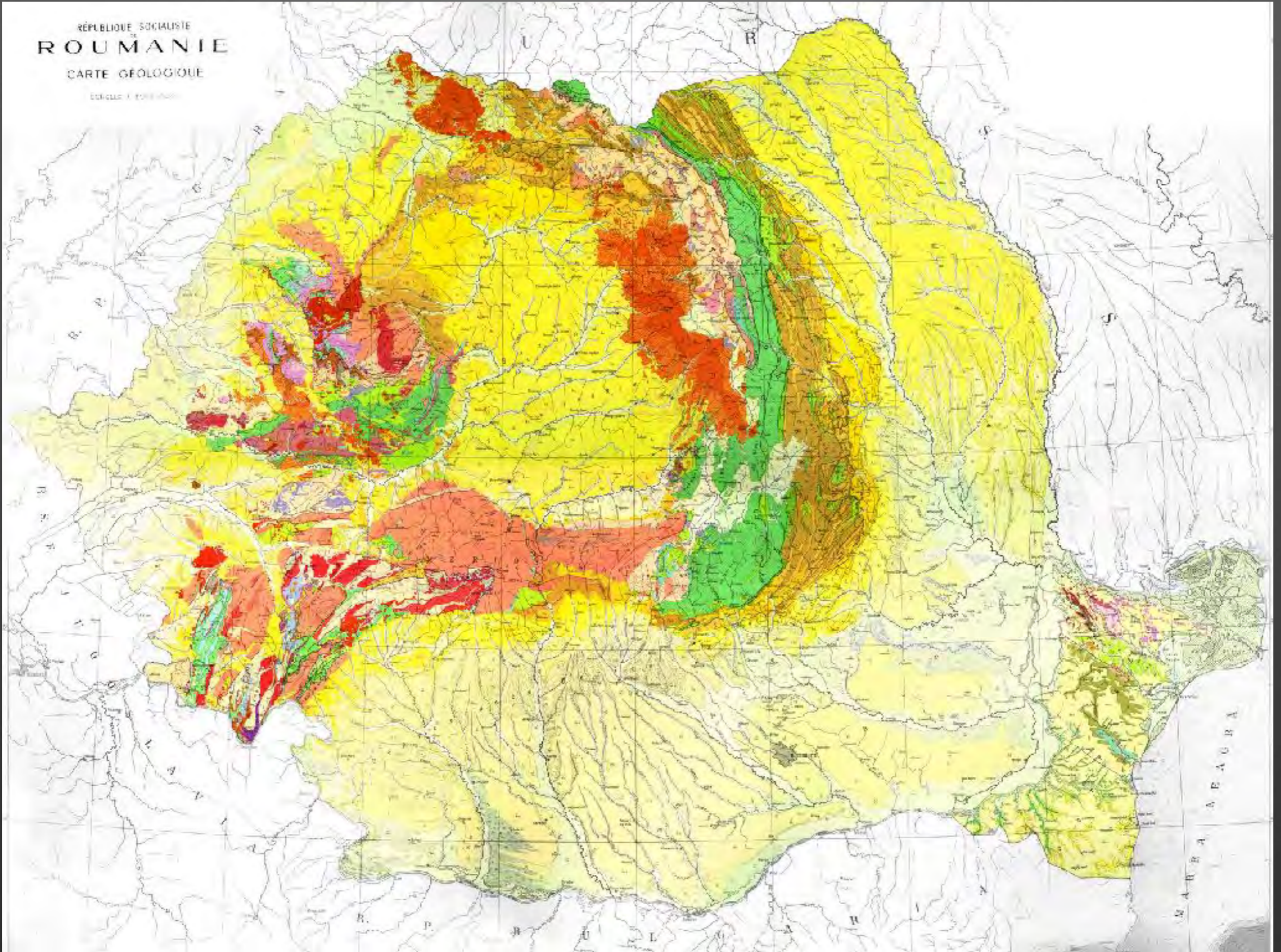


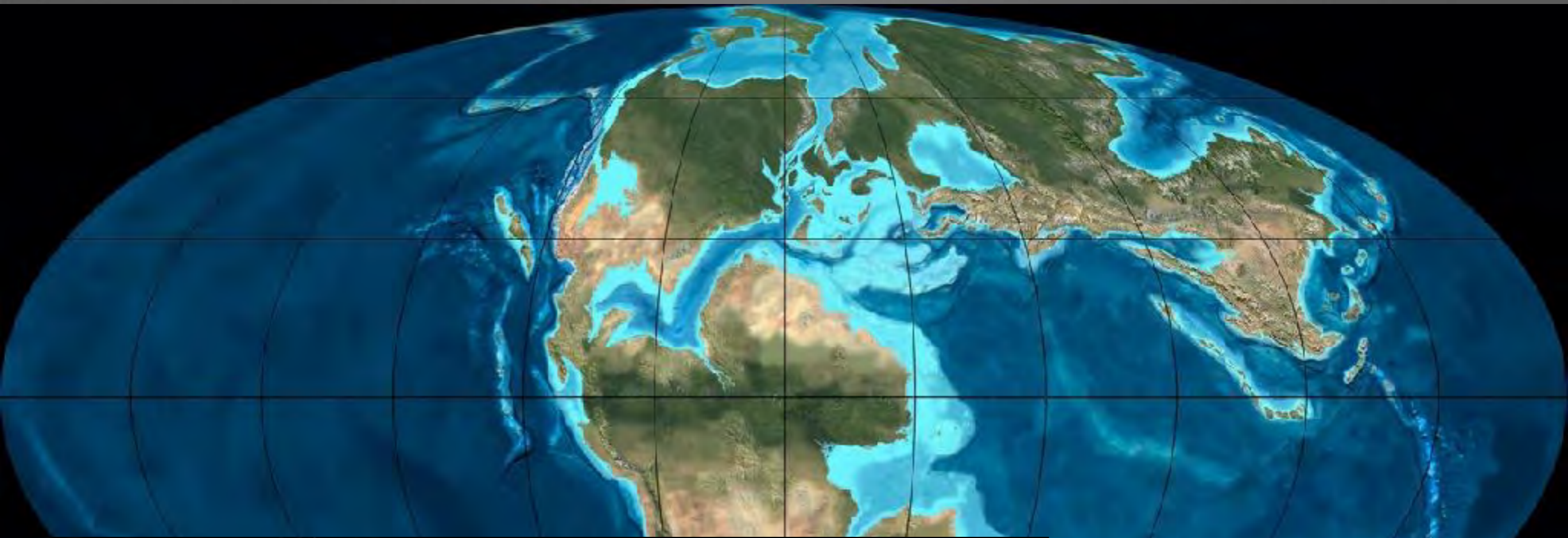
The Alpine story

- We were near but not at the real margin of the Tethys ocean (Mesozoic - to Cenozoic) and saw the development of back arc basins, maybe the main arc of the late Jurassic -early Cretaceous
- The back arc basins - a segment closed and made the South Carpathians, the rest evolved into a large marginal back arc basin known as the Paratethys
- The Paratethys closed in the East Carpathian (Miocene and younger) but continues to exist east in the Black Sea

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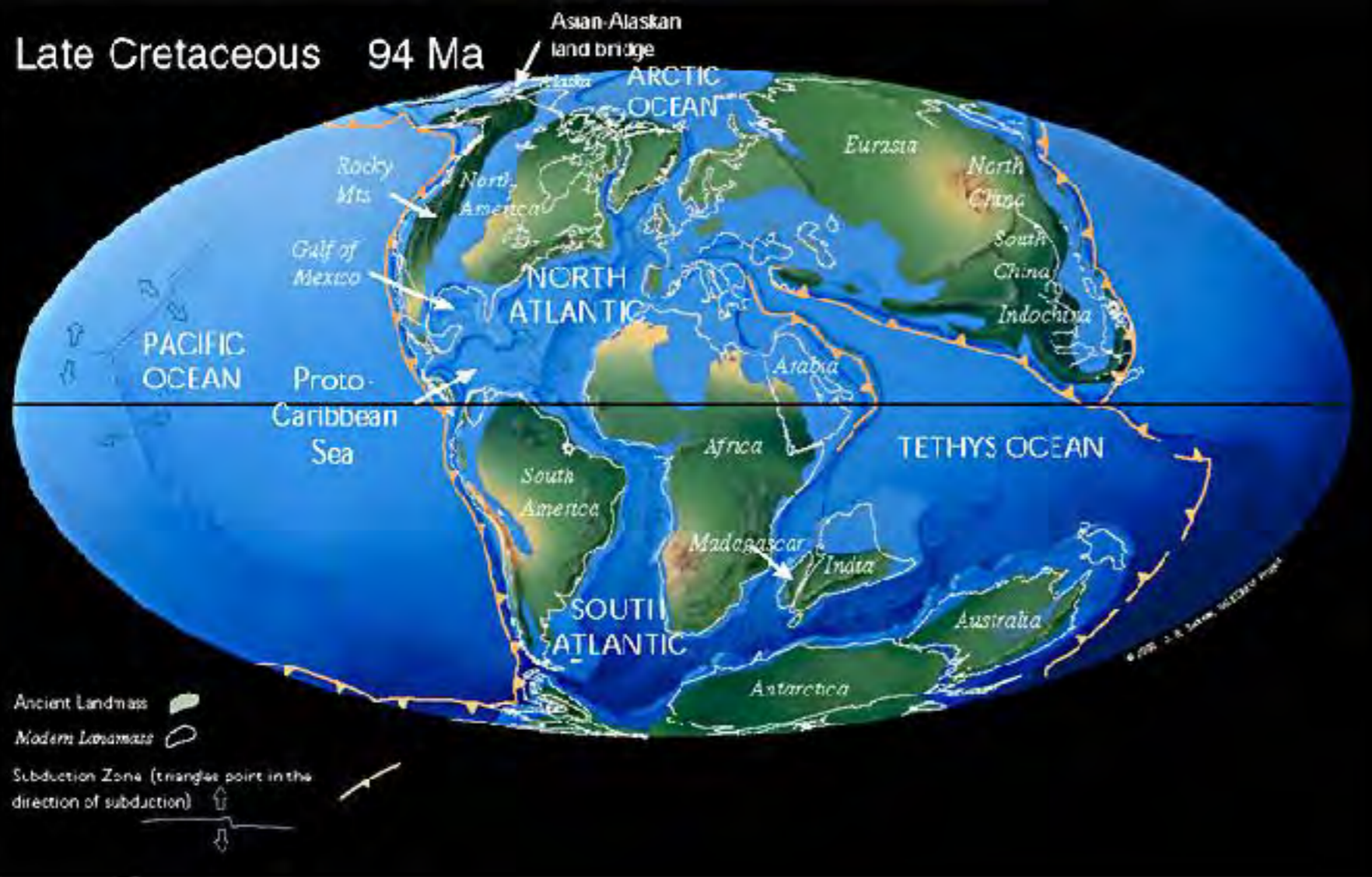
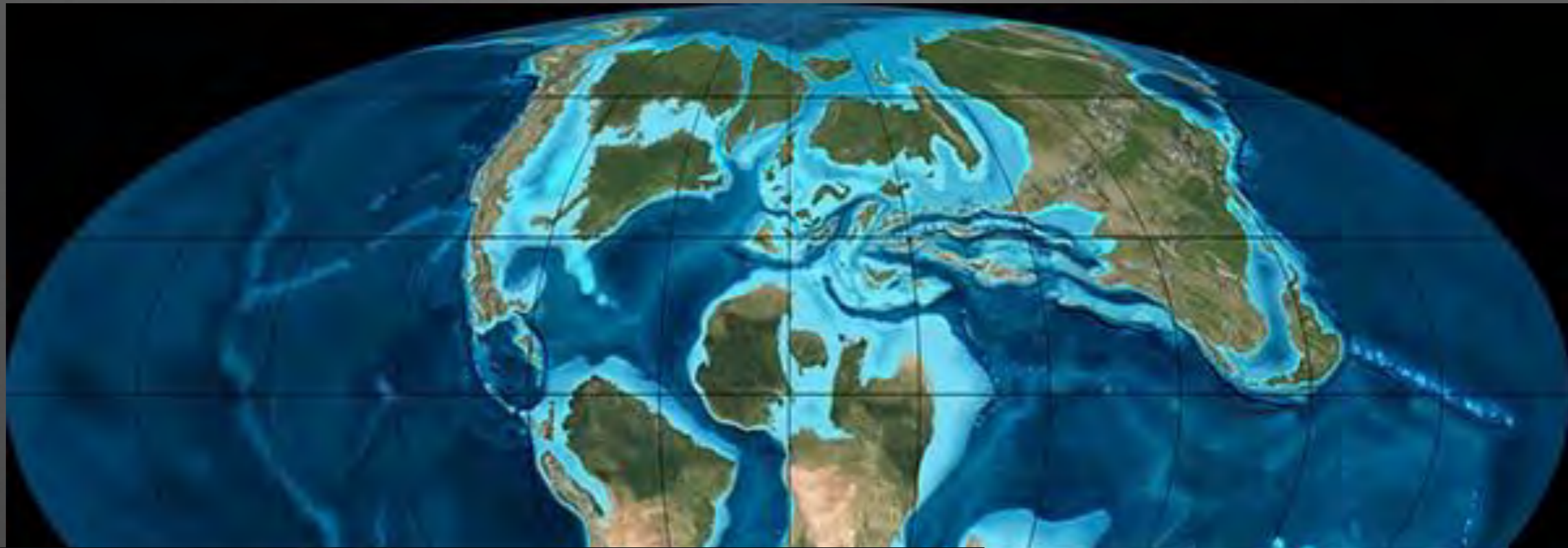


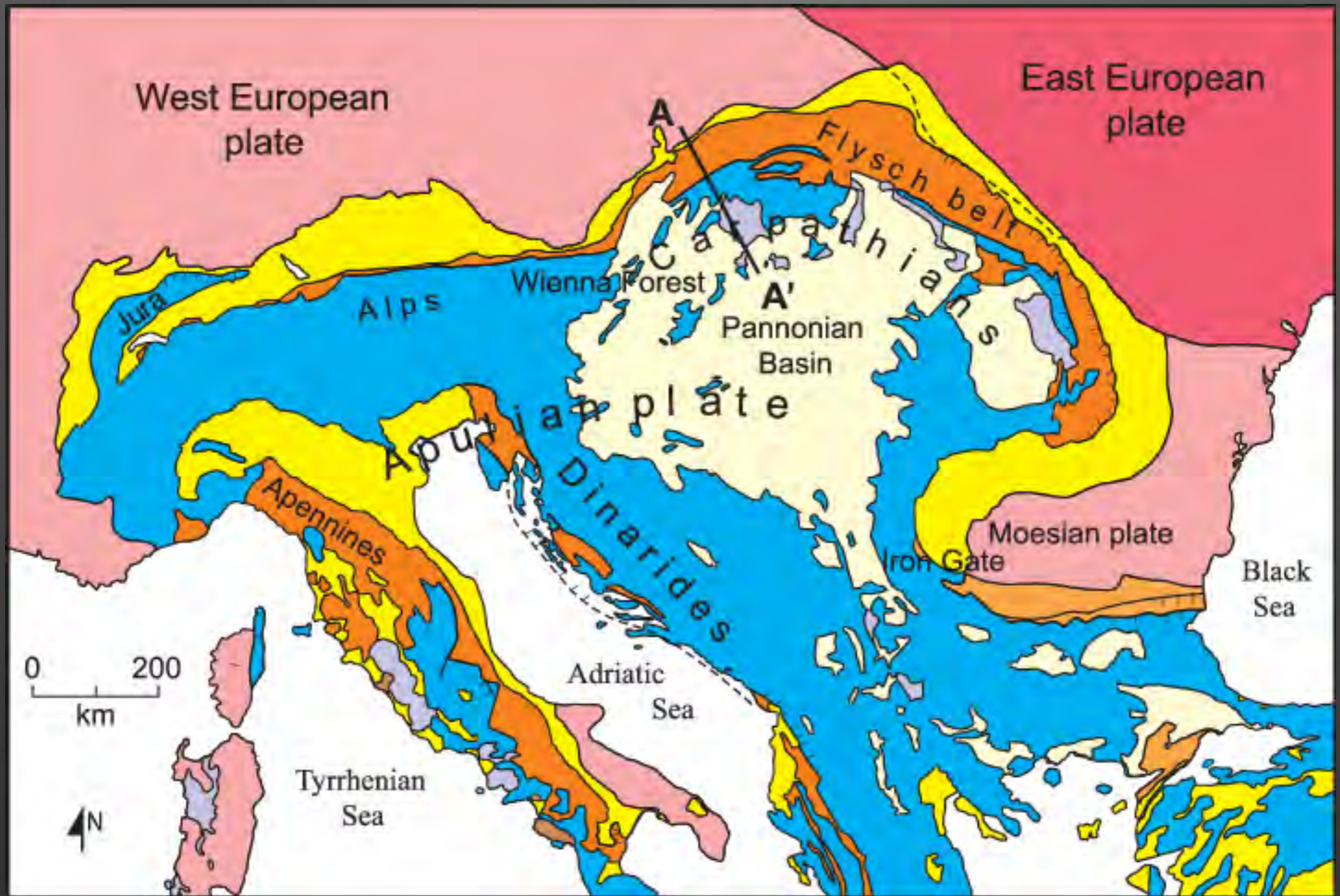


Late Jurassic 152 Ma

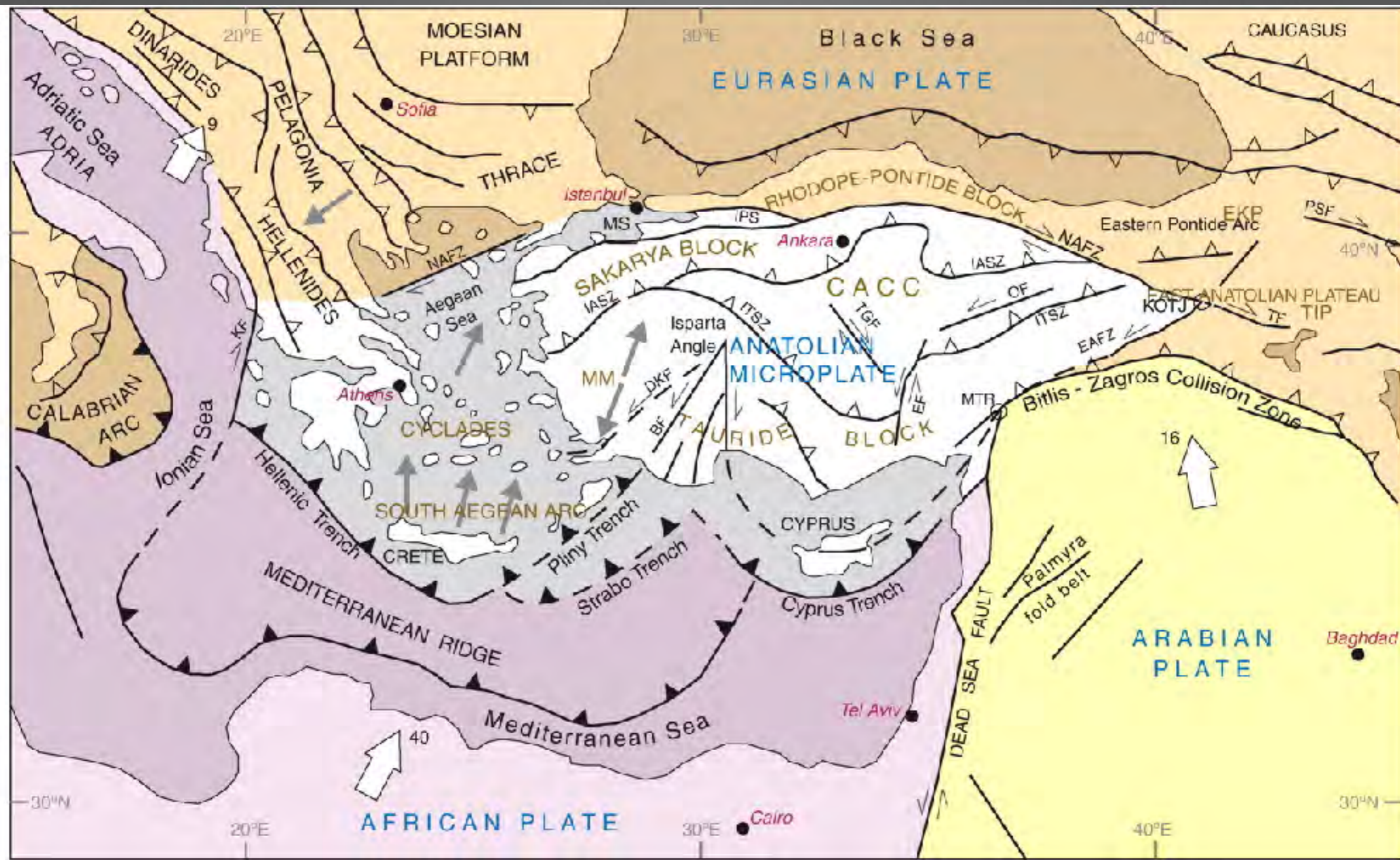


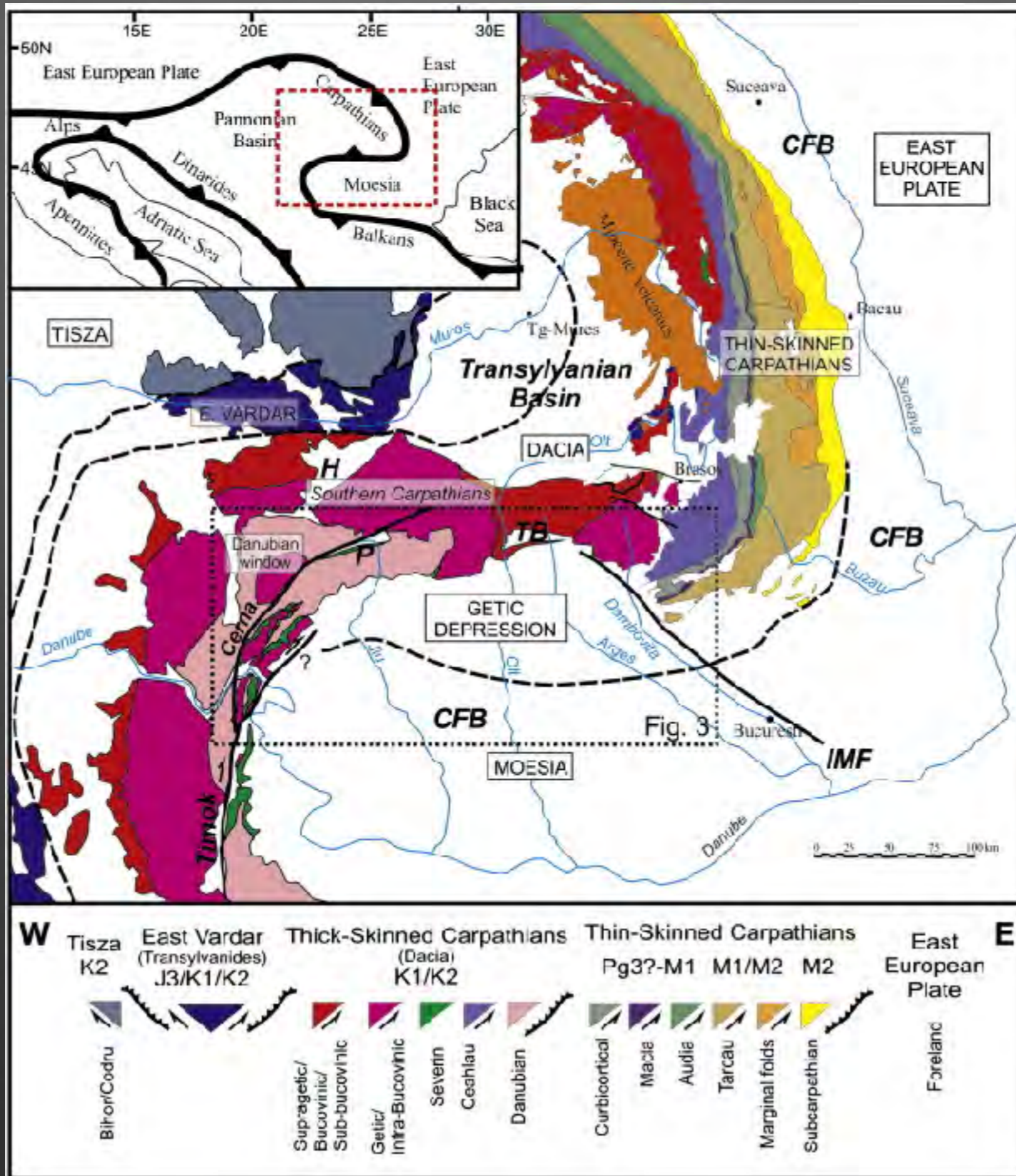
Ancient landmass 
 Modern landmass 
 Subduction zone
 (triangles point in the
 direction of subduction)
 Sea floor spreading ridge 

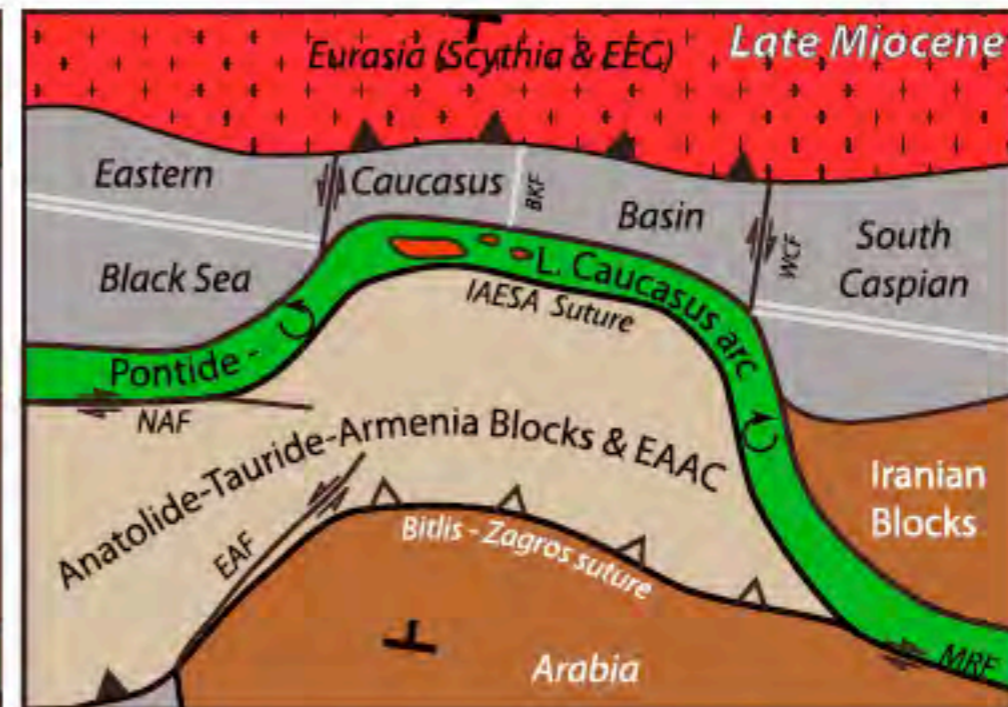
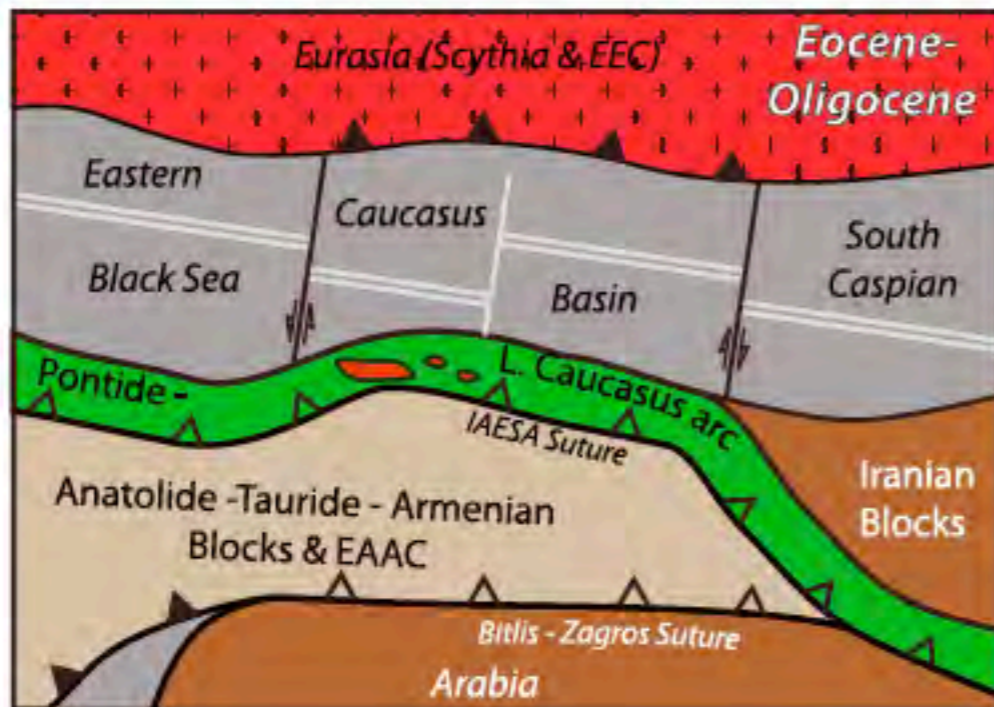
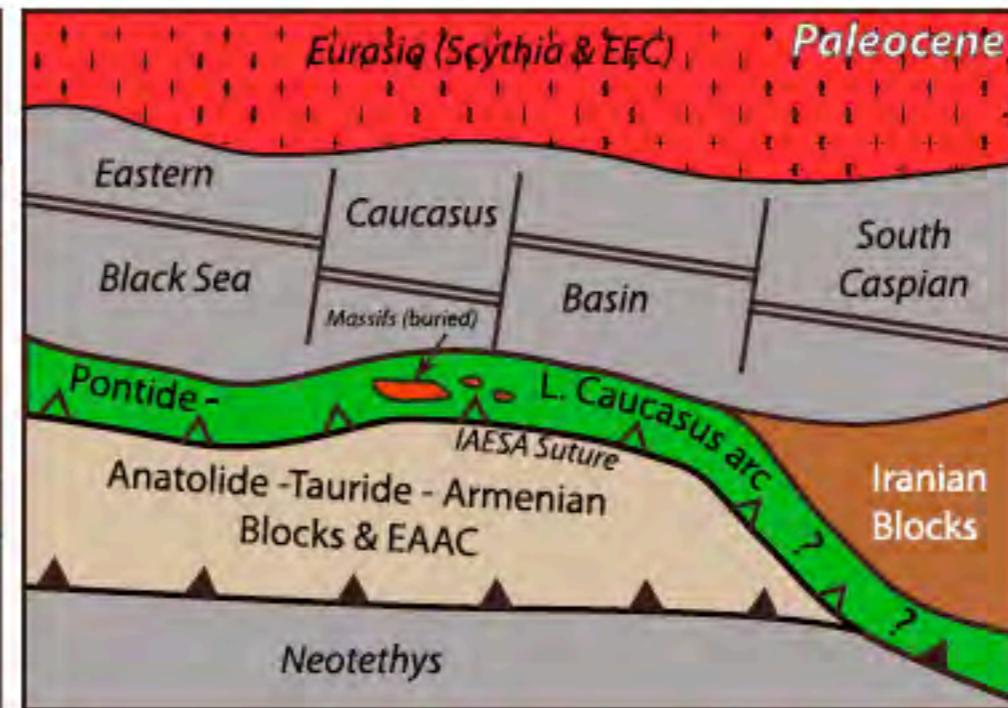
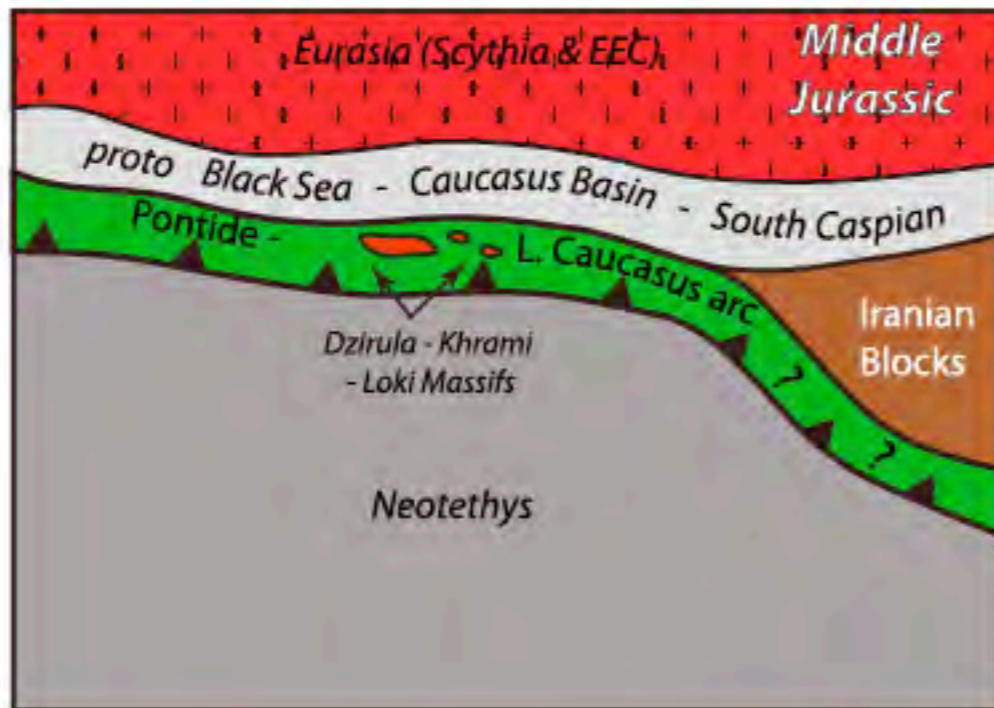


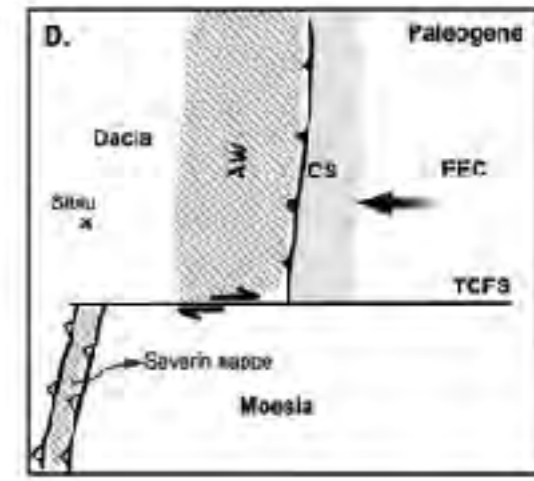
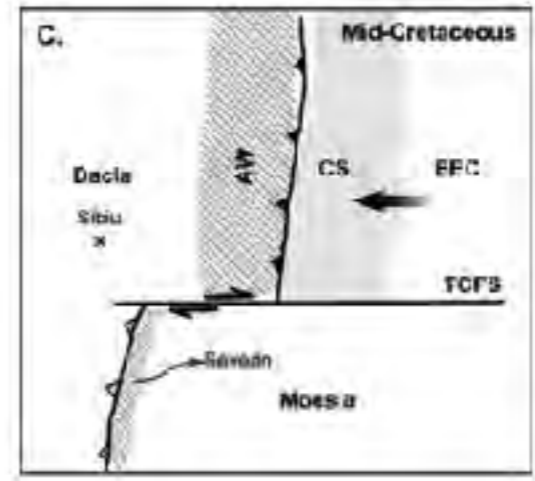
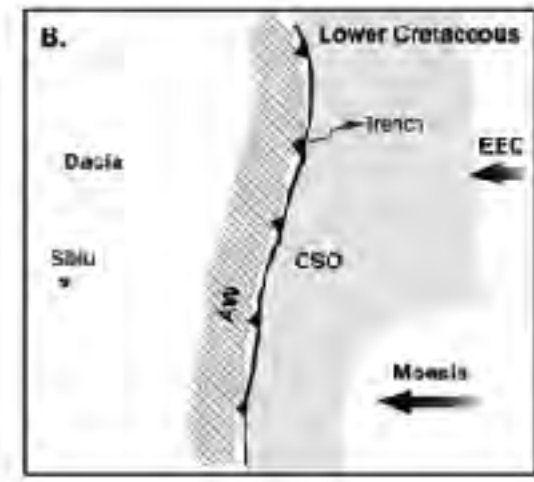
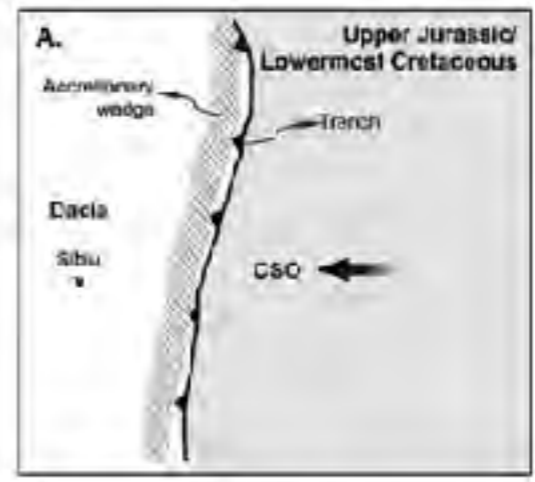
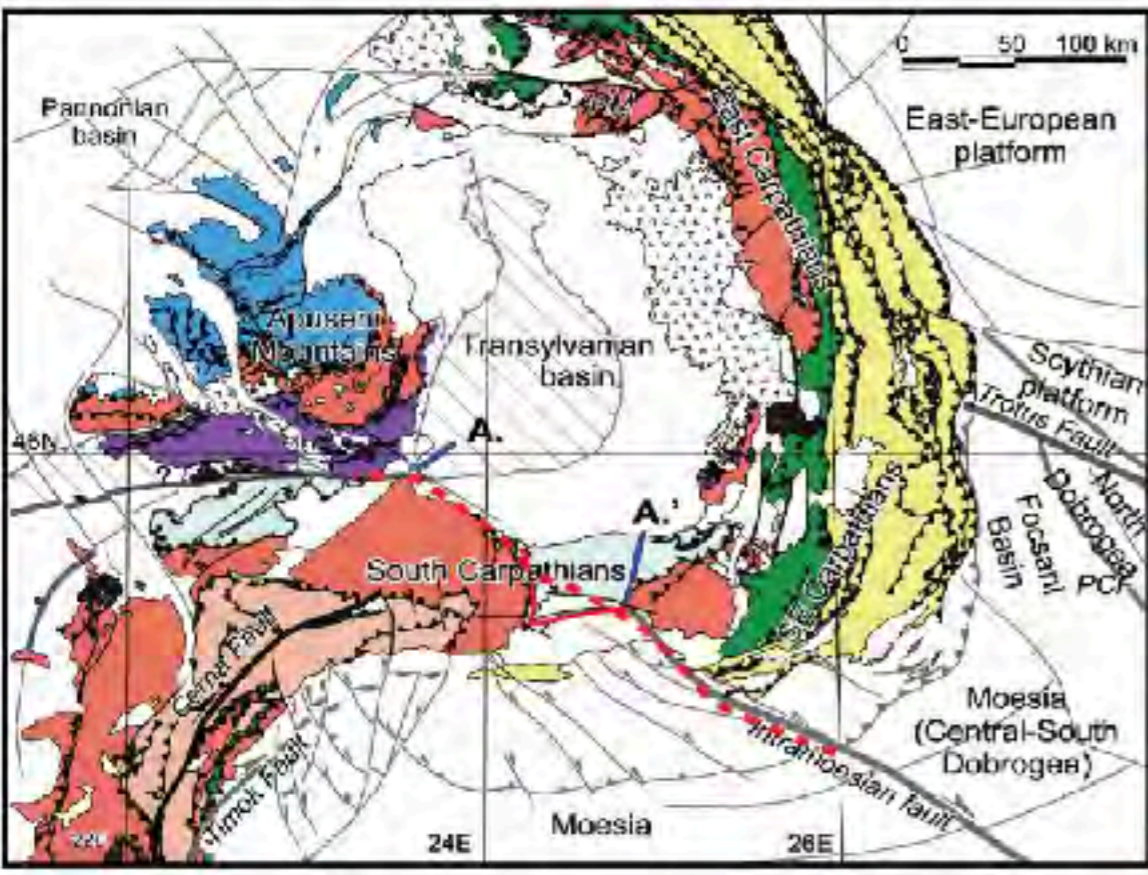


b)







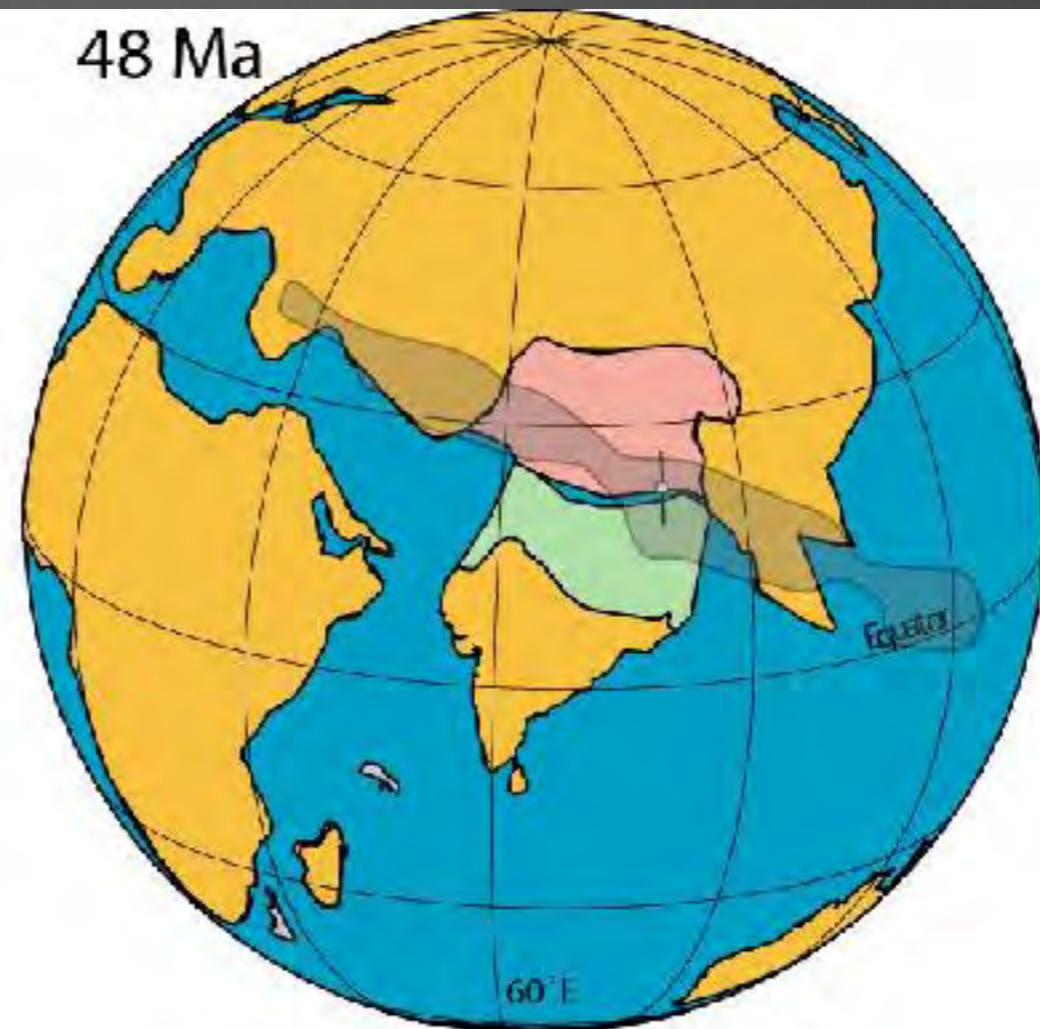
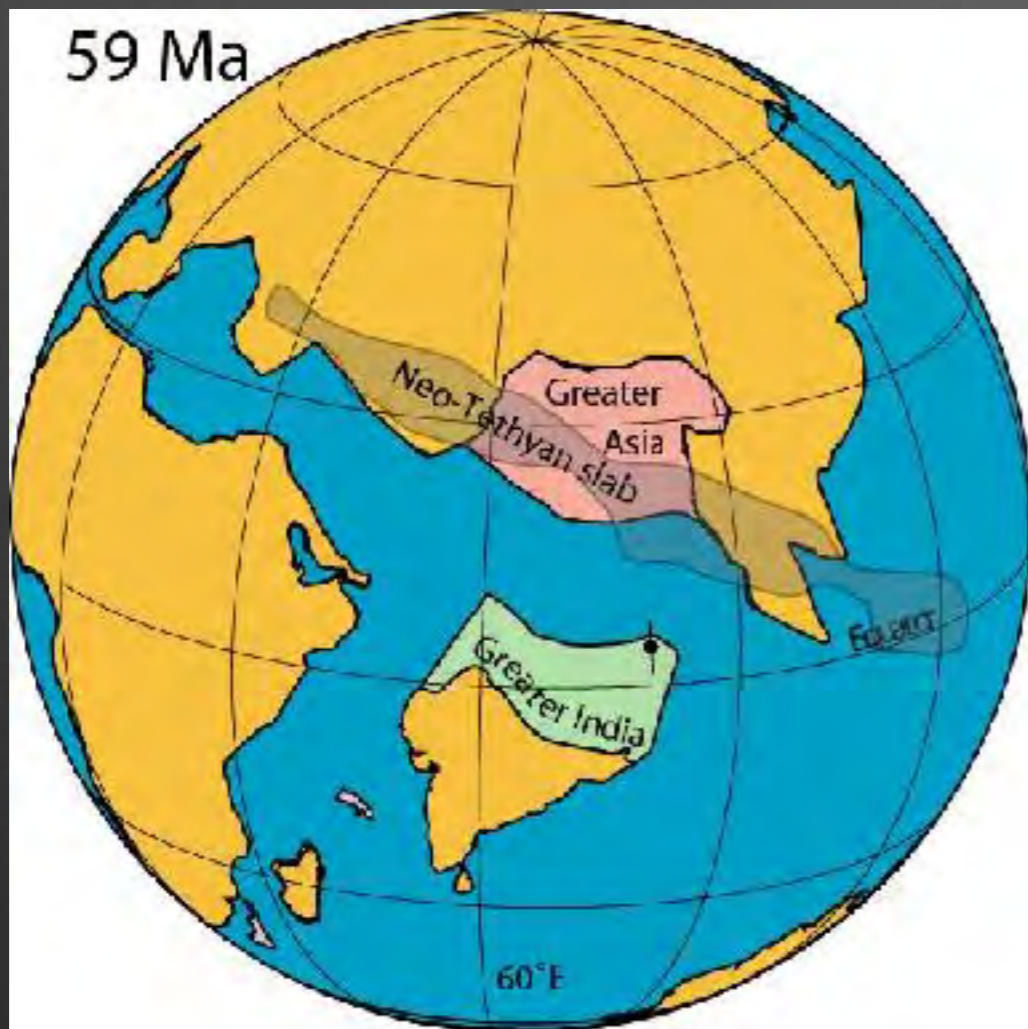


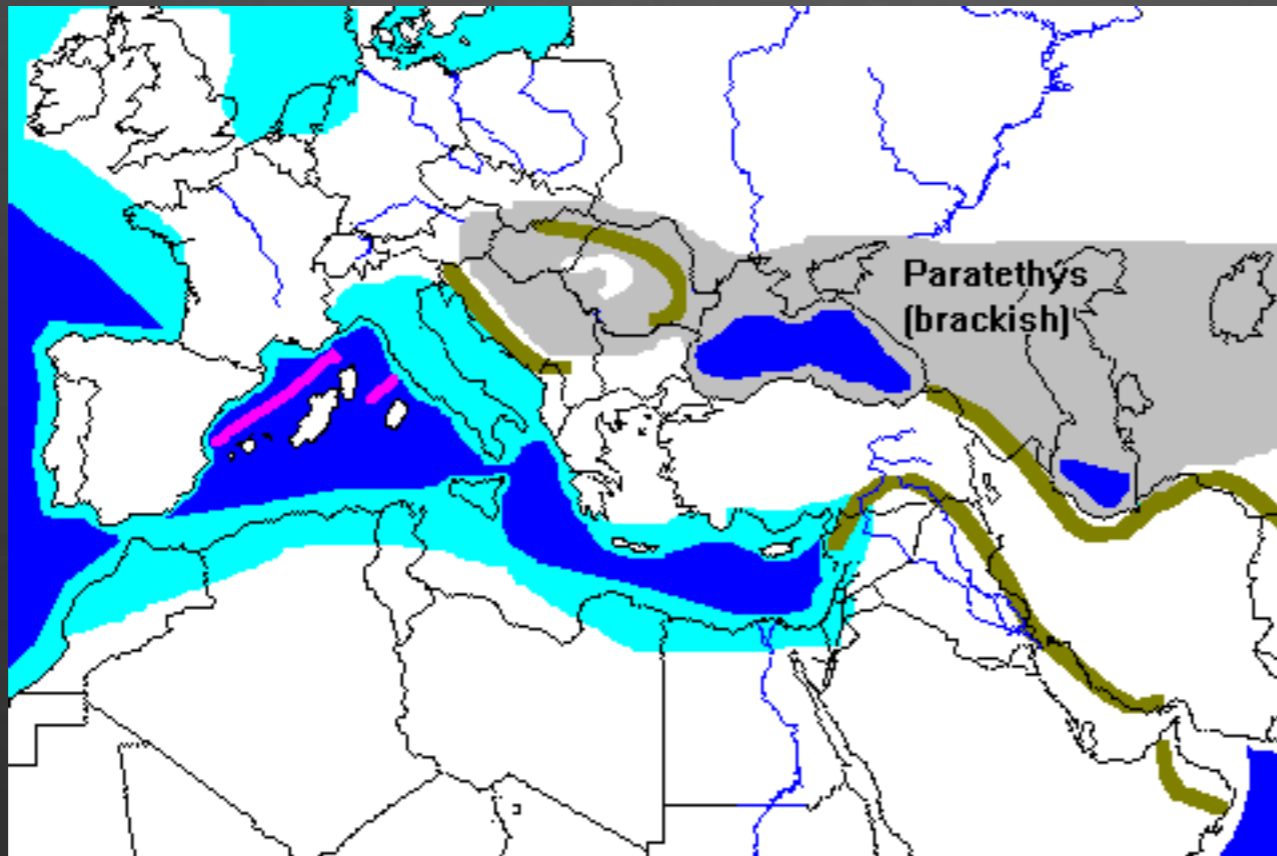
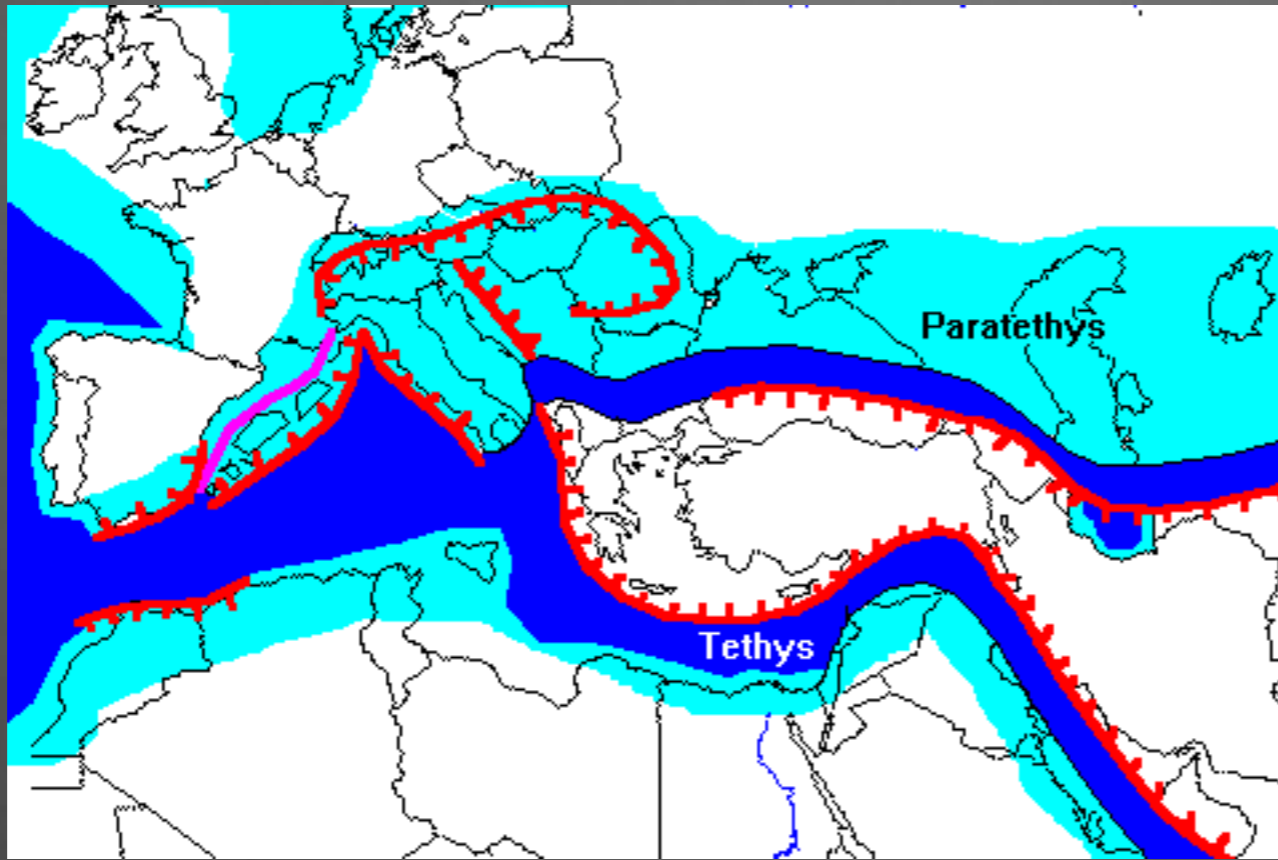
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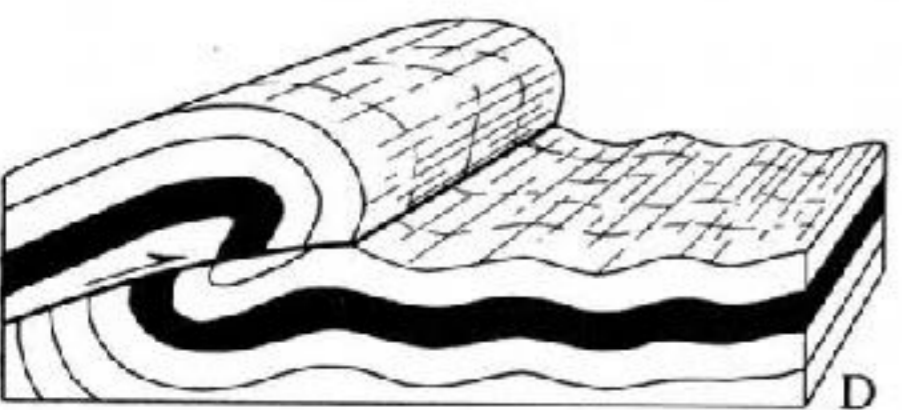
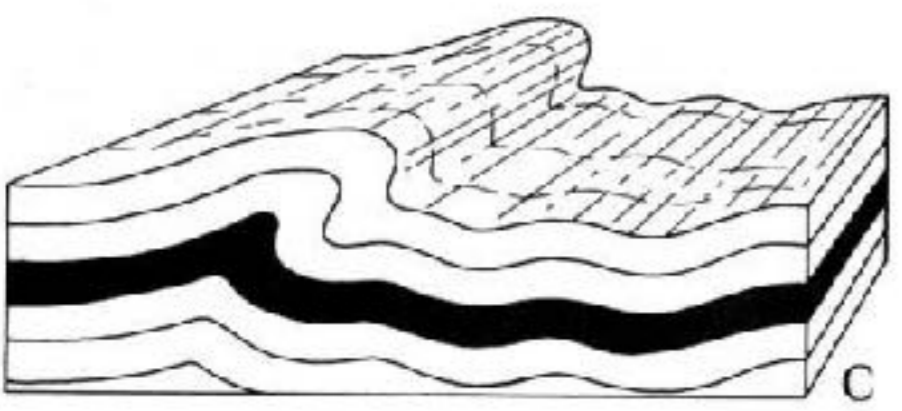
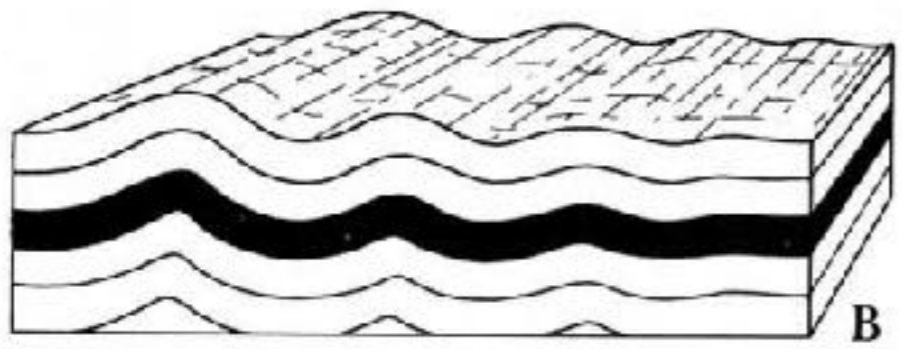
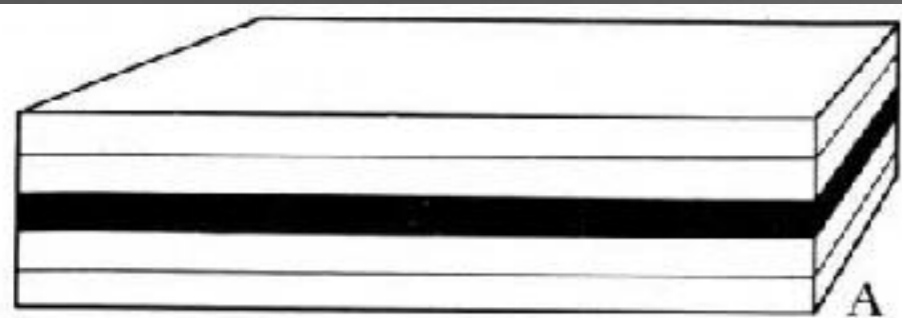
Figure 1:
 Reconstructed Eurasian palaeogeography in the late Eocene, 35 million years ago.
 (<http://cpgeosystems.com/>)
 This particular reconstruction shows an elongated, shallow Paratethys with open marine gateways to the eastern Tethys Ocean and North Sea.







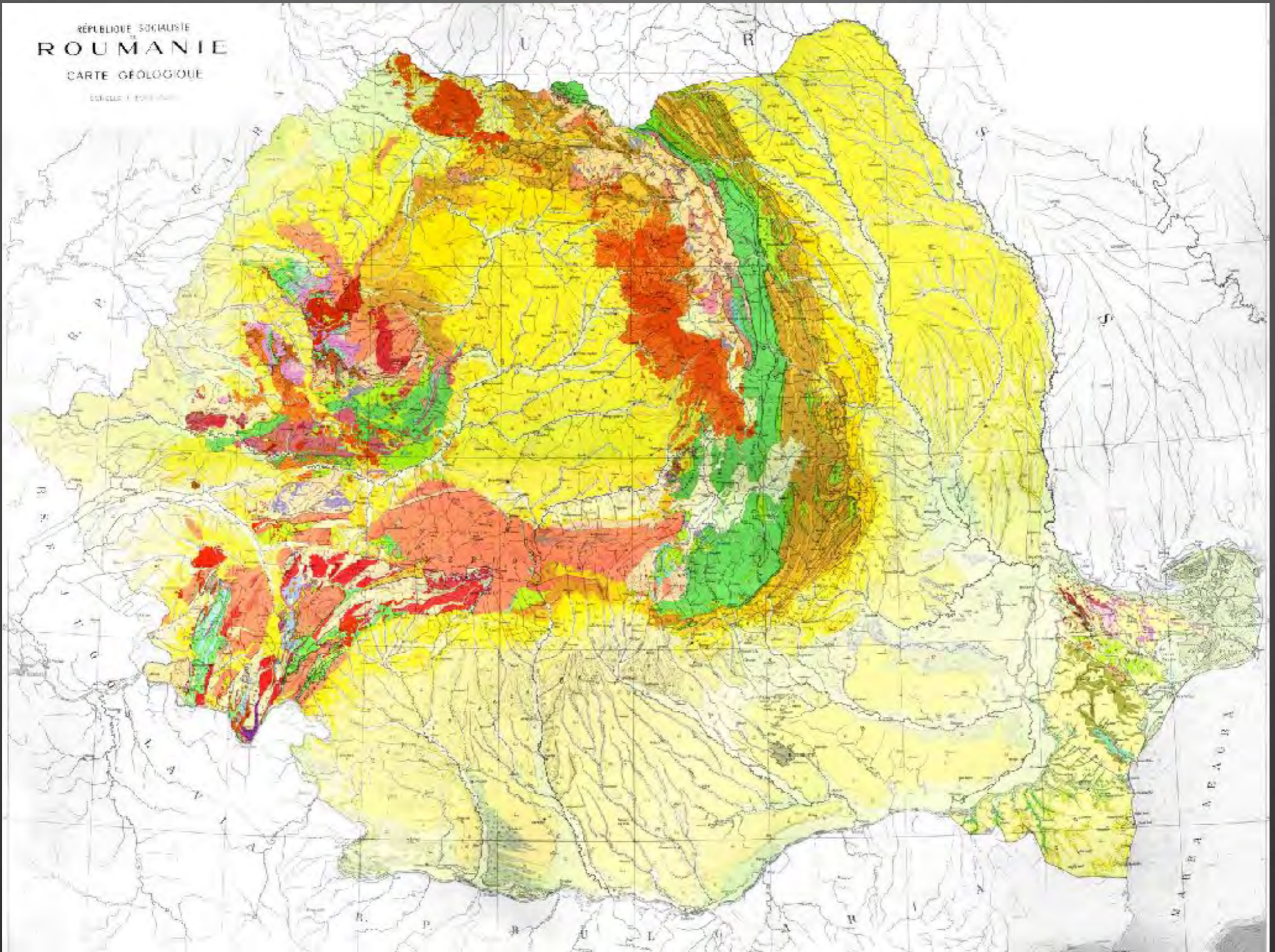
From: http://cpgeosystems.com/13_Mid_Mio_Eurmap_sm.jpg, July 2015

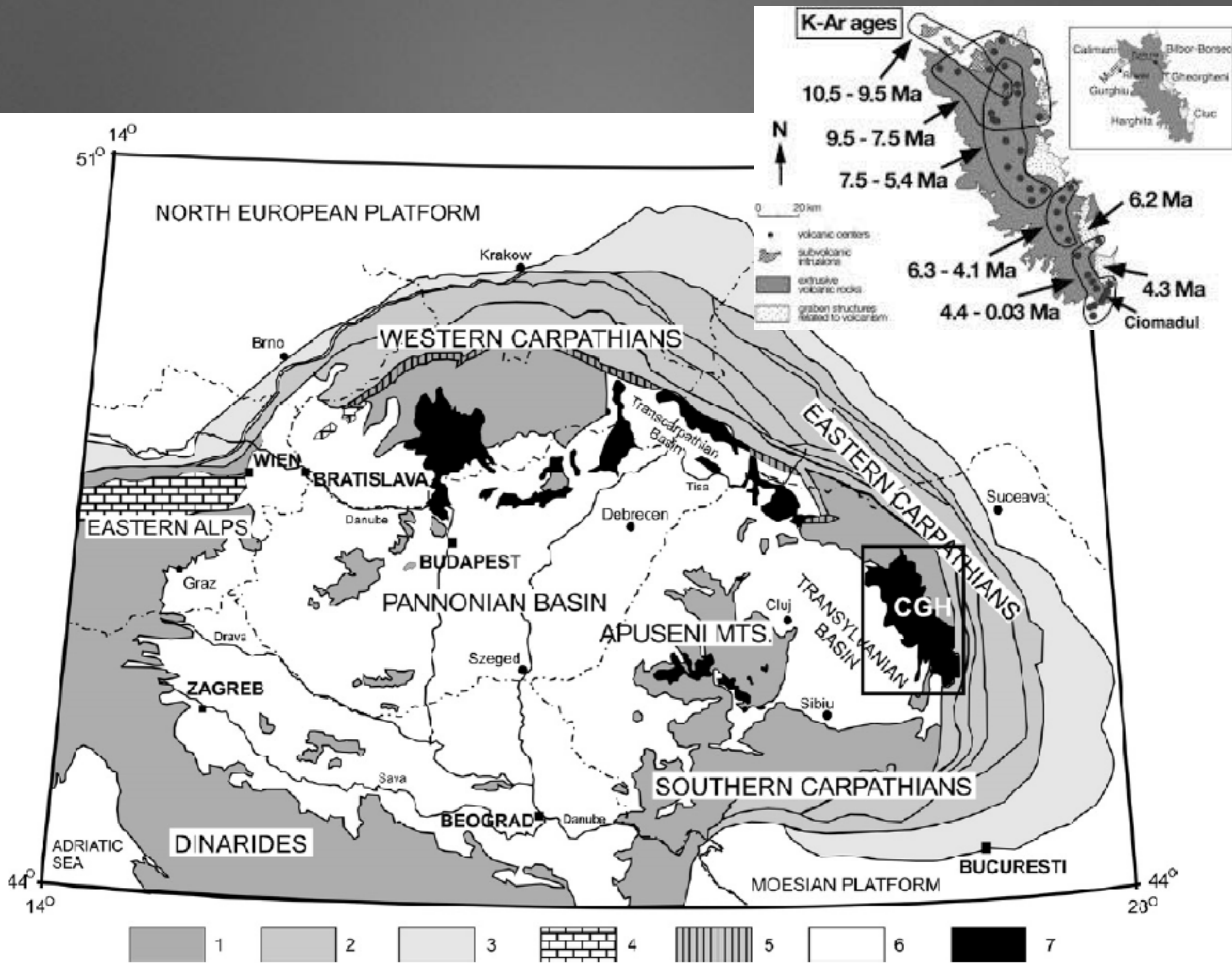


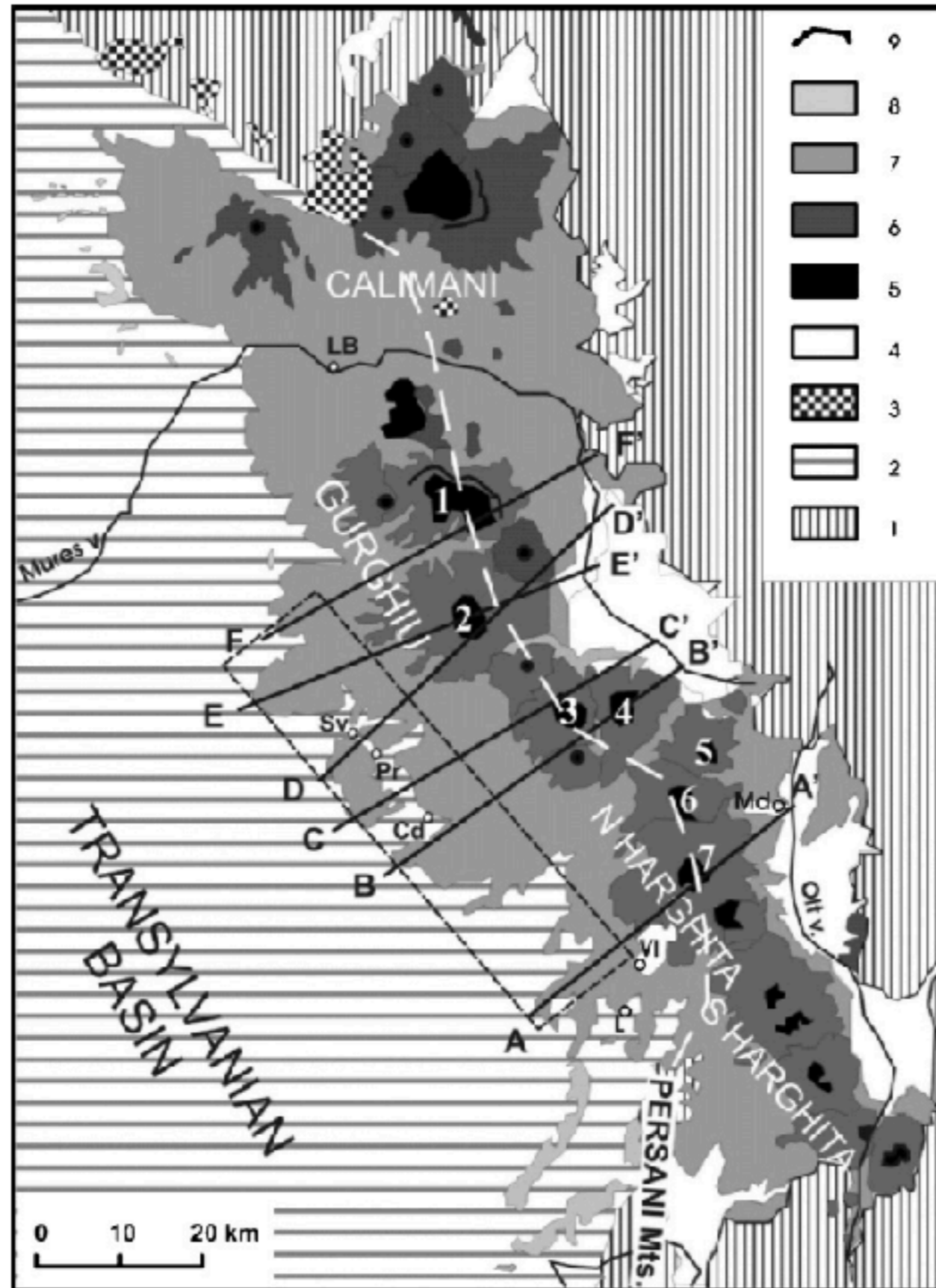
Volcanism

- Minor post basement formation
- Some Jurassic arc rocks in so Apuseni
- A belt of subduction rocks called “banatites”
- Neogene volcanism in Orientali and Apuseni

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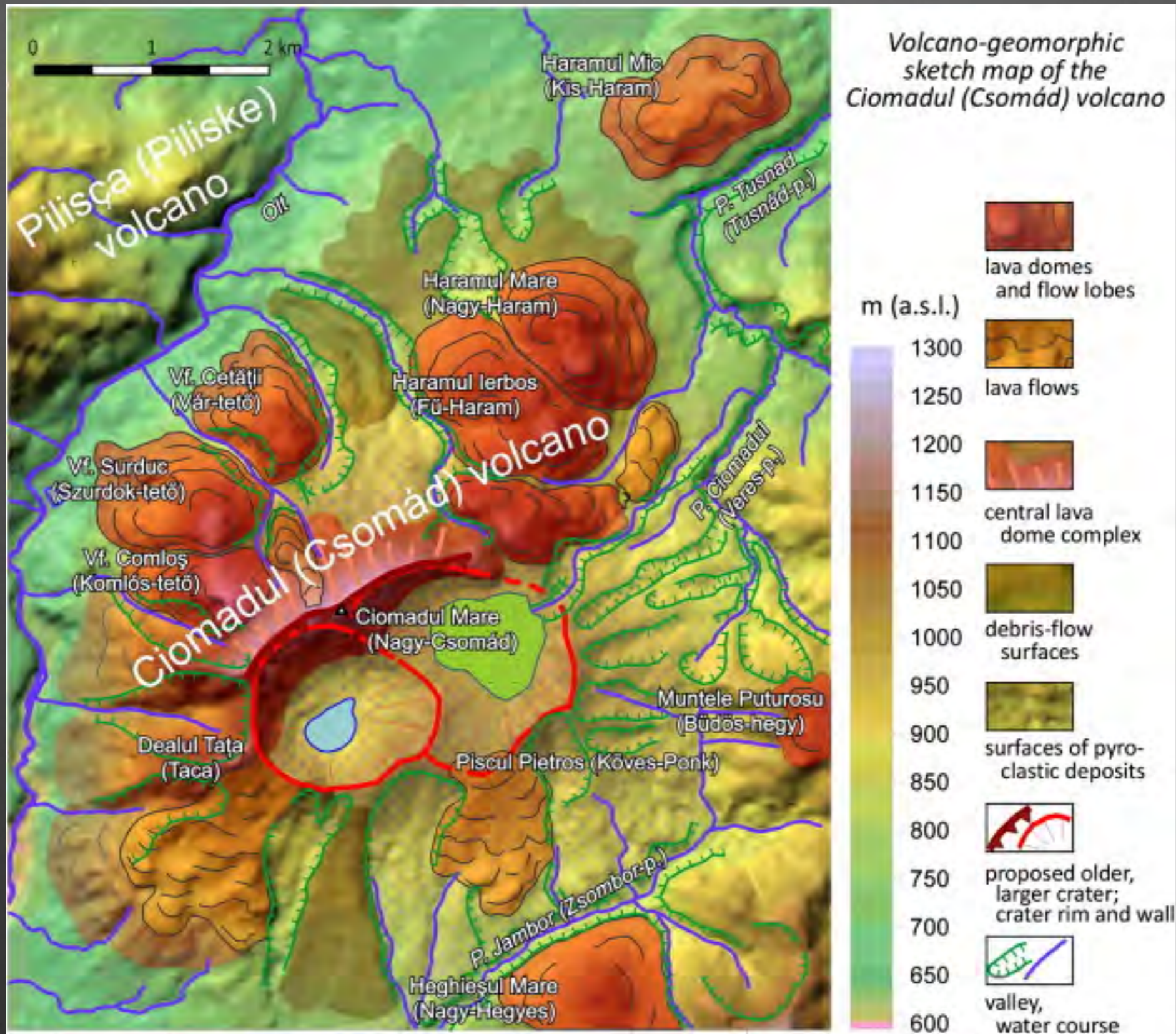


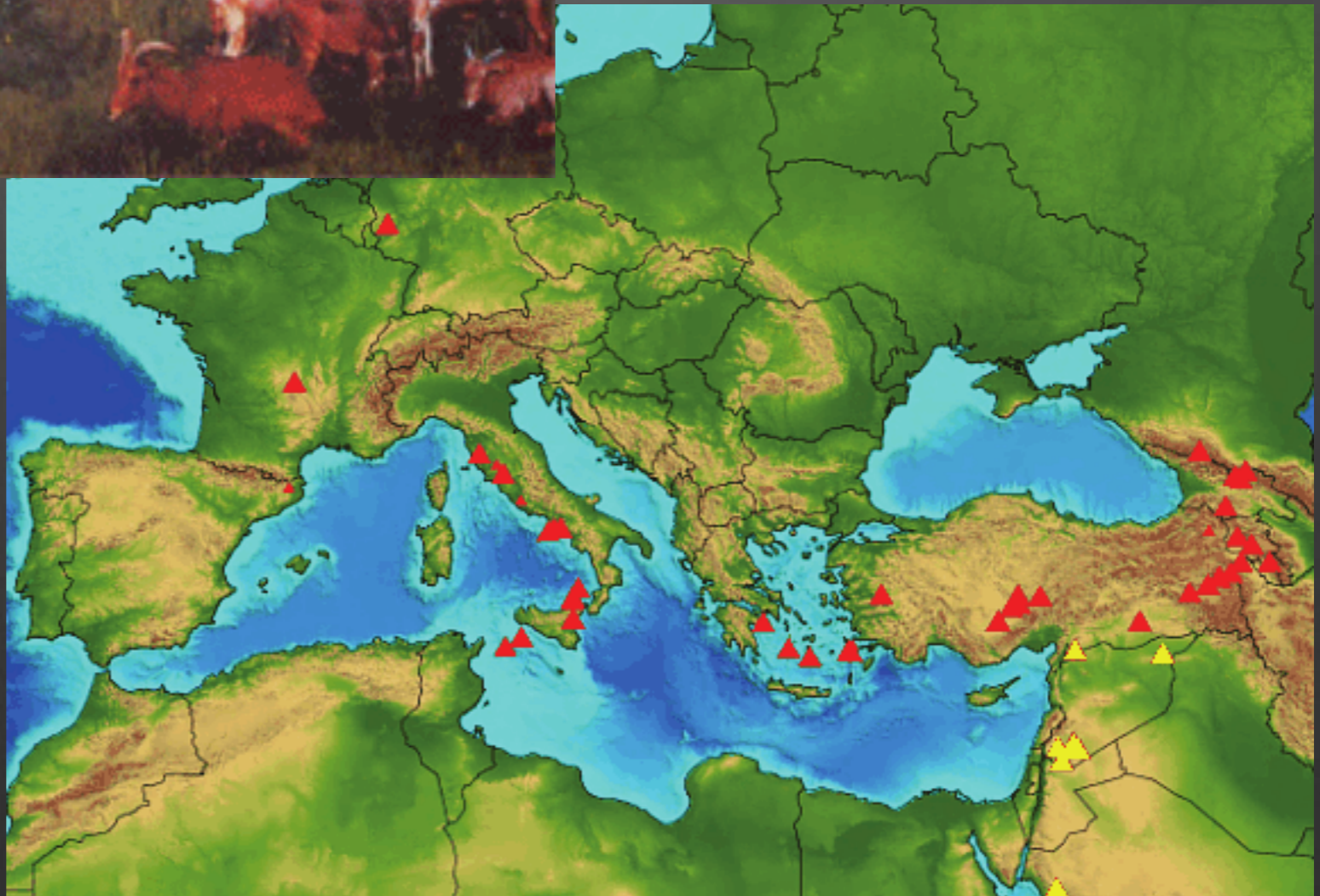


Origin of magmatism

- Compositions are subduction-related like but there is no obvious subduction -
- Ages progressing southward calls for some unusual geologic scenario
- They seem to parallel the timing of paratethys closure but how??/
- Youngest volcano is essentially active



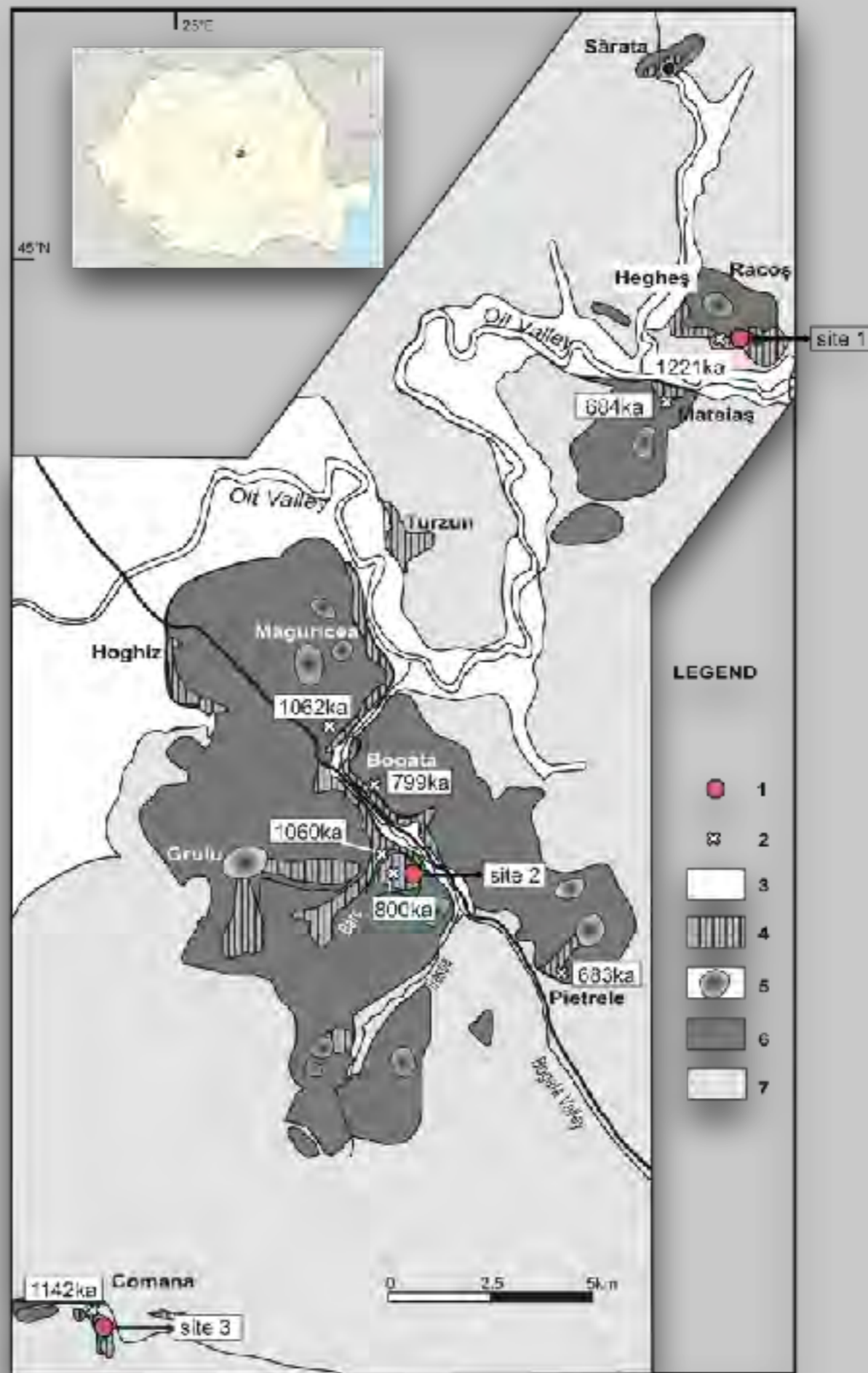




Persani

- Basaltic magmatism in a medium size magmatic field near Racos
- Mantle -derived
- Probably related to the Vrancea slab







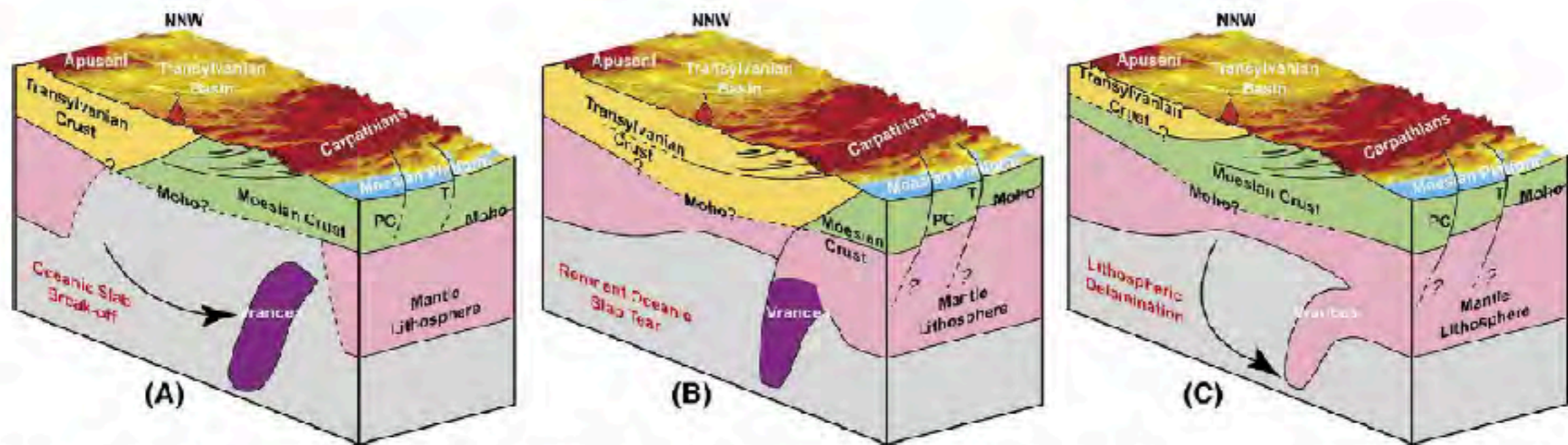
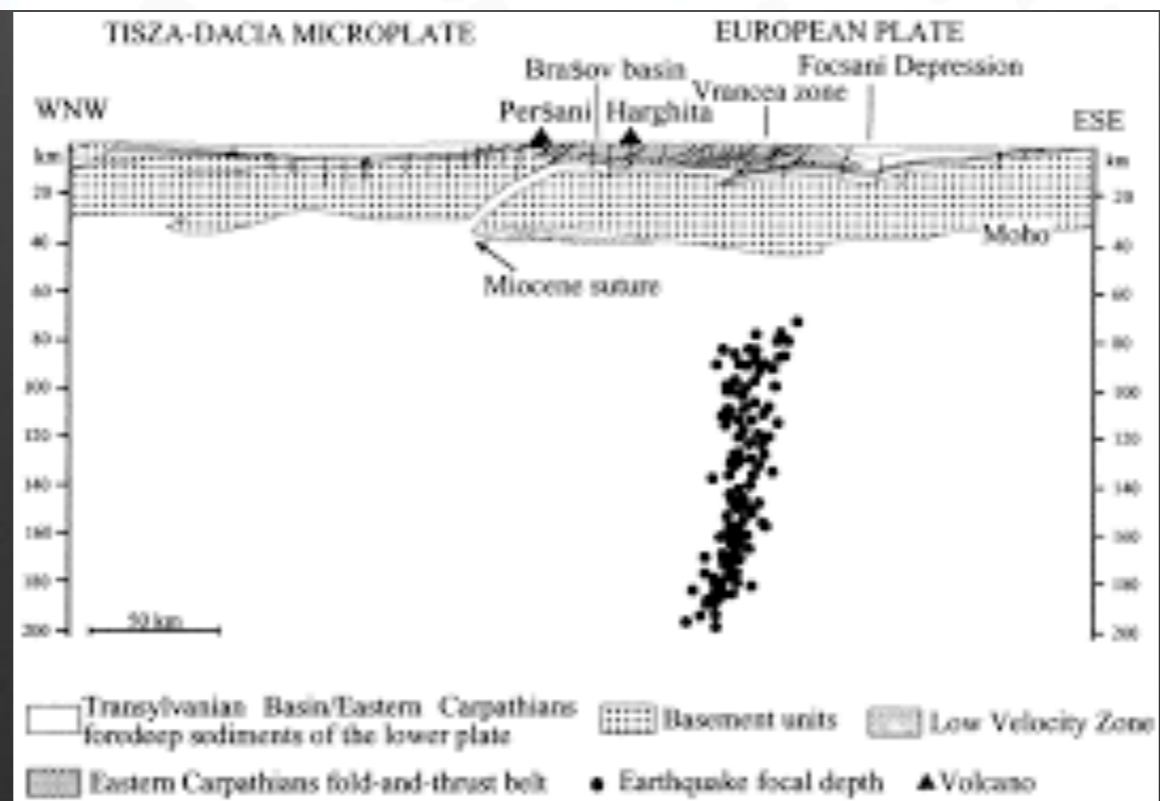


Fig. 6. 3-D perspective lithosphere-scale block models (view towards NNW), illustrating contrasting scenarios for geodynamic setting of the Vrancea zone. (A) Oceanic slab subduction and break-off. (B) Oceanic slab subduction and progressive lateral tear within the Carpathian foreland. (C) Continental lithospheric delamination. Green = Moesian/East European crust; yellow = Transylvanian crust; pink = continental mantle lithosphere; purple = oceanic lithosphere; grey = asthenosphere. Vrancea zone is located in lower front corner of models. See text for discussion.



Delamination?

- Roll back of lithosphere and detachment
- Makes way for mantle upwelling and melting
- Currently tested by our team











Photo by Jay Quade





